



A roadmap for amphibious drilling at the Campi Flegrei caldera: insights from a MagellanPlus workshop

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Received: 20 March 2019 - Revised: 14 September 2019 - Accepted: 26 September 2019 - Published: 2 December 2019

Abstract. Large calderas are among the Earth's major volcanic features. They are associated with large magma reservoirs and elevated geothermal gradients. Caldera-forming eruptions result from the withdrawal and collapse of the magma chambers and produce large-volume pyroclastic deposits and later-stage deformation related to post-caldera resurgence and volcanism. Unrest episodes are not always followed by an eruption; however, every eruption is preceded by unrest.

The Campi Flegrei caldera (CFc), located along the eastern Tyrrhenian coastline in southern Italy, is close to the densely populated area of Naples. It is one of the most dangerous volcanoes on Earth and represents a key example of an active, resurgent caldera. It has been traditionally interpreted as a nested caldera formed by collapses during the 100–200 km³ Campanian Ignimbrite (CI) eruption at \sim 39 ka and the 40 km³ eruption of the Neapolitan Yellow Tuff (NYT) at \sim 15 ka. Recent studies have suggested that the CI may instead have been fed by a fissure eruption from the Campanian Plain, north of Campi Flegrei.

A MagellanPlus workshop was held in Naples, Italy, on 25–28 February 2017 to explore the potential of the CFc as target for an amphibious drilling project within the International Ocean Discovery Program (IODP) and the International Continental Drilling Program (ICDP). It was agreed that Campi Flegrei is an ideal site to investigate the mechanisms of caldera formation and associated post-caldera dynamics and to analyze the still poorly understood interplay between hydrothermal and magmatic processes. A coordinated onshore–offshore drilling strategy has been developed to reconstruct the structure and evolution of Campi Flegrei and to investigate volcanic precursors by examining (a) the succession of volcanic and hydrothermal products and related processes, (b) the inner structure of the caldera resurgence, (c) the physical, chemical, and biological characteristics of the hydrothermal system and offshore sediments, and (d) the geological expression of the phreatic and hydromagmatic eruptions, hydrothermal degassing, sedimentary structures, and other records of these phenomena. The deployment of a multiparametric in situ monitoring system at depth will enable near-real-time tracking of changes in the magma reservoir and hydrothermal system.

1 Introduction

Large collapse calderas are associated with climactic explosive volcanic eruptions capable of producing a global catastrophe second only to that from a meteorite impact. On the other hand, many calderas are characterized by hydrothermal systems that represent a source of "clean", geothermal energy production (e.g., Lipman, 2000). Despite numerous scientific and applied studies, the inner structure and the dynamics of caldera systems are still poorly known. In many cases, large calderas also host densely populated urban and agricultural districts. As a consequence, understanding the caldera structure and dynamics has an immediate effect on the assessment of volcanic hazards and associated risk at a local and global scale.

Large calderas are found on all continents and in different geological settings. For example, recent restless examples can be found in New Zealand (Taupo), North America (Crater Lake, Long Valley, Valles, Newberry, and Yellowstone), South America (Laguna de Maule and Cerro Blanco), Asia and Oceania (Toba, Tambora, Krakatau, Rabaul, Toya, Shikotsu, and Kuttara), and Europe (Santorini and Campi Flegrei). Some of these are located close to coastlines and continental shelves, where hydrothermal and groundwater dynamics may partly control the expression of volcanism and the distribution of eruptive products. Deposition of these products occurs in a changing depositional regime with a high average sedimentary supply. Coastal and partly submerged calderas on continental shelves thus contain unique stratigraphic archives of interbedded volcaniclastic and marine deposits with a high potential for preservation.

The Campi Flegrei caldera (CFc), next to Naples in southern Italy, has been the world's most restless, non-erupting caldera for the last 69 years, characterized by episodes of significant ground uplift, enhanced hydrothermal activity, and seismicity. In addition to these short-term episodes (e.g., 1950–1952, 1969–1972, and 1982–1984; De Natale et al., 2006; Del Gaudio et al., 2010; Troiano et al., 2011; Chiodini et al., 2015, 2017; Kilburn et al., 2017; Moretti et al., 2017, 2018), ground subsidence and uplift of several meters has been documented since at least Roman times (e.g., Bellucci et al., 2006; Di Vito et al., 2016). Moreover, as a result of resurgence over the last ca. 12 000 years, the central part of the CFc has undergone a long-term, antiformal uplift of ca. 100 m that is partly recorded by La Starza marine terrace's present-day elevation (ca. 40 m above sea level) that emerged ca. 5000 years ago (Di Vito et al., 1999; Sacchi et al., 2014).

The cause of ground uplift episodes and phases occurring at the CFc (magmatic vs. hydrothermal) is still debated and likely consists of periods of shallow magmatic intrusions accompanied by injections of deep fluids into shallow aquifers (e.g., De Vivo and Lima, 2006; De Natale et al., 2006; Lima et al., 2009; Moretti et al., 2018; Troise et al., 2019). Significant hydrothermal activity is shown on land and in the submerged portion of the caldera by the discharge of hot gases and liquids (Sacchi et al., 2014; Passaro et al., 2016; Chiodini et al., 2017; Steinmann et al., 2018). Since offshore emissions cover an area 4 times larger than the main onshore hydrothermal site around the Solfatara crater (Passaro et al., 2014; Somma et al., 2016; Steinmann et al., 2018), the marine portion of the CFc may play a substantial role in the recent dynamics of the caldera, representing an underestimated source of degassing and heat flux. The lack of adequate data on offshore fluid emissions prevents a correct estimate of the fluid release and gas and heat flow budget at Campi Flegrei. In addition, while the uppermost 100 m of the submerged part of the CFc have been intensively studied (D'Argenio et al., 2004; Milia and Torrente, 2007; Sacchi et al., 2009, 2014; Passaro et al., 2016; Steinmann et al., 2016, 2018), the deeper portion remains largely unknown.

Understanding the mechanisms for unrest and eruptions is of primary importance for confident hazard assessment. Data on the deeper, submerged portion of Campi Flegrei are required to constrain forecasts of the type, intensity, and frequency of future magmatic, phreatomagmatic, and hydrothermal eruptions. More than 600 000 people are potentially exposed to pyroclastic flows, rising to about 2 million considering the ash fall, also emitted from submerged vents (Rossano et al., 2004; Mastrolorenzo et al., 2006, 2008; Tonini et al., 2015; Sandri et al., 2016, 2018). Traditionally, calderas have been analyzed through field studies, monitoring observations, analogue models, and numerical simulations (e.g., Druitt and Sparks, 1984; De Natale et al., 1991, 2001, 2006; Martí et al., 1994, 2008; Gudmundsson et al., 1997; Gudmundsson, 1998; Burov and Guillou-Frottier, 1999; Acocella et al., 2000, 2001, 2004; Martí and Gudmundsson, 2000; Roche et al., 2000; Roche and Druitt, 2001; Folch and Martí, 2004; Lavallée et al., 2004; Geyer et al., 2006; Gregg et al., 2012, 2013). More recently, offshore reflection seismic imaging has emerged as a tool to understand the stratigraphic architecture and shallow structure of collapse-resurgent calderas in continental margins (e.g., Sacchi et al., 2009, 2014; Passaro et al., 2016; Steinmann et al., 2016, 2018). However, only deep drilling can provide conclusive information on the causes and mechanics of unrest and on the state and evolution of the magmatic-hydrothermal system. These data represent a fundamental prerequisite for evaluating the caldera-related hazards (Lowenstern et al., 2017).

2 Campi Flegrei caldera

Campi Flegrei is an active caldera belonging to the Neapolitan Volcanic District, which includes the active volcanoes of Vesuvius and Ischia Island. The caldera contains the westernmost districts of Naples as well as the towns of Pozzuoli, Bacoli, Baia, and Quarto and several smaller villages. Half of the CFc is submerged and forms Pozzuoli Bay (also known as the Gulf of Pozzuoli). This area has represented an active segment of the eastern Tyrrhenian margin since the Late Quaternary (Oldow et al., 1993; Ferranti et al., 1996) and may be considered a natural laboratory for studying the interplay between tectonics and explosive volcanism in the rifted backarc margin of the Tyrrhenian Sea and the Adriatic subduction system below the Apennine fold-and-thrust belt (Fig. 1) (e.g., Milia and Torrente, 1999; Acocella et al., 1999).

The CFc describes a quasi-circular depression approximately 13 km across. The present-day shape of the caldera has been conventionally interpreted as the result of two large collapses related to the eruptions of the Campanian Ignimbrite (CI, ~ 39 ka; Giaccio et al., 2017) and the Neapolitan Yellow Tuff (NYT, ~ 15 ka; Deino et al., 2004) (Fig. 2) with respective volumes of 200 km³ DRE (dense-rock equivalent) (Rolandi et al., 2003) and 40 km³ DRE (Scarpati et al., 1993). Evidence of older ignimbrites has been reported in the Campanian Plain (De Vivo et al., 2001) and in the distal marine archives (e.g., Insinga et al., 2014). The locations of these eruptions remain poorly constrained around Campi Flegrei. As described below, an alternative view now emerging is that the CI was erupted outside Campi Flegrei so that the caldera was formed only by the NYT eruption.

The CI eruption is Europe's largest explosive volcanic event recorded in the last 200 000 years. It has been considered a possible cause for the decline of the Neanderthals, thus implying a potential influence on human evolution (Fitzsimmons et al., 2013). The CI deposits are widespread in the Mediterranean, and its ash has been found in the Russian Plain, more than 2500 km away from the source (Pyle et al., 2006; Giaccio et al., 2008). The CI eruption was followed by the NYT eruption at 15 ka and at least 60 post-caldera eruptions (Di Vito et al., 1999). The most recent eruption occurred in 1538 after a repose of ca. 3000 years. It produced Monte Nuovo and was preceded by a century of uplift (Bellucci et al., 2006; Di Vito et al., 2016). Most recently, noneruptive unrest episodes occurred during 1950-1952, 1969-1972, and 1982–1984 (Del Gaudio et al., 2010). They have been characterized by ground deformation (with rates up to $100 \,\mathrm{cm} \,\mathrm{yr}^{-1}$), shallow, low magnitude earthquakes (about 16 000 events with a magnitude up to 4.0 in 1983–1984), and marked geochemical variations in the emitted gases (Berrino et al., 1984; De Natale and Zollo, 1986; Dvorak and Berrino, 1991; De Natale et al., 1991, 1995, 2001; Battaglia et al., 2006; Chiodini et al., 2015; Di Luccio et al., 2015; Moretti et al., 2017, 2018). In fact, the recorded history of non-eruptive ground movements goes back to Roman times as revealed by marine incrustations and mollusks on Roman and medieval buildings (e.g., Bellucci et al., 2006; Troise et al., 2019). Campi Flegrei thus represents the caldera with the longest record of ground movements not immediately followed by eruptions. Uplift began again in 2005 after 20 years of subsidence. It has been characterized by a slow movement of ca. 3 cm yr^{-1} and less seismicity but a longer duration than previous uplifts (Troise et al., 2007; Chiodini et al., 2017; Moretti et al., 2018). The long duration of the ongoing unrest has led the Civil Protection Department to declare the first level on its alert scale (yellow), which implies an increase in monitoring activities.

Although the CI eruption has been previously considered as the main caldera-forming event (Rosi and Sbrana, 1987; Orsi et al., 1996), De Vivo et al. (2001) and Rolandi et al. (2003) presented evidence in favor of a fissure eruption to the north of the CFc. Recently, new evidence of buried CI products inside the caldera area at a depth of ca. 400 m beneath the surface has been found in the pilot borehole of the ICDP (International Continental Drilling Program) Campi Flegrei Deep Drilling Project (CFDDP) (De Natale et al., 2016). The shallow depth and modest thickness of the deposit (less than 200 m) raised further questions about a caldera collapse associated with the CI eruption. These and other studies highlight the complexity of the caldera system and, hence, the need for additional in situ information to fully understand the whole framework and evolution of volcanism in Campania. Due to its partly submerged setting, Campi Flegrei represents an ideal site to test the potential of IODP (International Ocean Discovery Program) shallow-water drilling on a volcanic continental margin by a multiplatform drilling pro-

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Figure 1. Tectonic sketch map of the Campanian continental margin in SW Italy with the location of the Campanian volcanic zone. Frame of Fig. 2 is also indicated.

gram including a land-sea transect, in the frame of a fully integrated IODP-ICDP Amphibious Drilling Proposal (ADP). The research outcomes derived from Campi Flegrei may also be transferred to similar partly submerged volcanic systems, including the Aira, Kikai, Krakatoa, Maug, Santorini, and Tavui calderas.

3 The MagellanPlus workshop

During the MagellanPlus workshop held in Naples on 25-28 February 2017, 35 participants from four European countries (Italy, Germany, Spain, and the UK), the USA, and Japan, gathered to discuss the key scientific issues for a coordinated IODP-ICDP proposal dedicated to drilling in the CFc. The workshop built upon previous research and networking activities, including (1) a coordinated ICDP and ESF (European Science Foundation) Magellan workshop held on 13-15 November 2006 in Naples; (2) an approved ICDP full proposal (Campi Flegrei Deep Drilling Project -CFDDP) in 2006–2008; (3) a submitted IODP pre-proposal (#671) in 2006–2007 with an indication of resubmission on the basis of an implemented site-survey package; (4) the realization of two pilot holes, a few meters apart, 500 and 200 m deep, as a preliminary phase to the ICDP deep drilling (the 200 m hole has been continuously cored by wireline drilling and used to install a borehole seismometer); and (5) the acquisition of new offshore site-survey data (3-D multiscale multichannel and single-channel reflection seismics, multibeam bathymetry, and gravity core data) between 2008 and 2016 (Fig. 3).

Participants at the MagellanPlus workshop represented a wide range of disciplines, including volcanology, geology, geophysics, geomorphology, petrology, geochemistry, and geochronology, as well as numerical and analogue modeling. The initiative was intended to bring together experts, early-career researchers, and other representatives from academia and industry involved in marine and continental research drilling. The aims were to (1) provide a global perspective on the potential and challenges of scientific drilling at active calderas, (2) discuss drilling issues on volcanic continental margin settings, (3) illustrate the new site-survey data, and (4) define drilling objectives for reconstructing stratigraphic events associated with the caldera's evolution and the interaction between magmatism and hydrothermal activity in coastal marine settings.

Participants were asked to contribute to scientific debates on volcanism and associated hazards over coastal areas and identify problems that can be addressed by coordinated marine and continental drillings, with reference to the CFc as a representative case study. The workshop program addressed data integration, the building of a scientific rationale for drilling strategies, and scientific partnering through a multidisciplinary approach, by linking geology, geophysics, volcanology, petrology, microbiology, and geotechnology. The event is among the first efforts to assess scientific themes directly related to volcanic hazards in highly populated coastal



Figure 2. Digital Terrain Model (DTM) of the Campi Flegrei caldera (CFc) showing the structural border(s) and the offshore ring-fault–resurgent-dome system associated with the eruption of the Neapolitan Yellow Tuff (NYT) (~ 15 ka), along with the most prominent subaerial and submerged volcanic morphologies of Napoli Bay. Location of seismic reflection profiles illustrated in Figs. 10 and 11 is also indicated.

areas within the context of fully integrated IODP-ICDP drilling research.

Fundamental questions were discussed on a wide range of topics, such as the mechanisms and timing of caldera formation and resurgence, ignimbrite deposition environments, magma transfer processes, and explosive volcanic activity in submarine and coastal settings, volcano-tectonic coupling, the dynamics and energy budget of onshore and offshore hydrothermal (geothermal) systems, subaerial vs. submarine volcanic unrest, and monitoring. Participants identified the following key questions and objectives that, which shall be addressed by the Amphibious Drilling Proposal:

- Interaction between magmatic and hydrothermal processes. This will be investigated regarding shallow crustal levels, the mechanism of submarine degassing and hot fluid discharge and their contribution to deformation, and recent unrest. What are the source, dynamics, and consequences of the hydrothermal activity in the marine portion of Campi Flegrei, and how are they related to unrest? Does the structural framework of the CFc exert control on the ascent of fluids and magma? Can microbial communities help in tracing hydrothermal fluid paths and defining thermodynamic environments and facies at depth?

- Stratigraphy and structures of the CFc. This will be investigated within the half-graben system of the Bay of Naples. This investigation includes recovering a representative stratigraphic record of the caldera fill and borders, down to the upper structural levels of the caldera floor and reconstructing the distribution of the CI (and older ignimbritic) deposits across the Bay.
- Kinematic reconstruction of caldera collapse structure and resurgence. This investigation includes reconstructing the pattern, timing, and rates of deformation involving the various structural components of the CFc system.
- *Eruptive history of the CFc*. When was the onset of volcanic activity, and what was the driving mechanism,



Figure 3. Offshore site-survey seismic data package (high-resolution multichannel and single-channel reflection seismic profiles) supporting the IODP component of the Campi Flegrei Amphibious Drilling Proposal. Locations of the proposed IODP drill sites and the ICDP site and pilot hole drilled in 2012 are also indicated.

e.g., volcano–earthquake interaction? In what way did the magmatic system change over time? Are there any major pre-CI ignimbrites originating from the Campanian Plain? What is the type of post-collapse, submarine volcanic activity? How have large eruptions changed the submarine morphology, and which are the syn- and post-eruptive deposition mechanisms?

- Difference between terrestrial and shallow marine volcanism. Why have post-collapse eruptions occurred preferentially in the onshore part of the caldera? Is the hydrothermal activity in the marine setting driving overpressure and fluid-saturated subvolcanic intrusions that are substantially different from on-land subvolcanic or volcanic processes?
- Environmental and climatic impact. This investigation refers to large-scale ignimbrite eruptions such as the CI and NYT events. Did these eruptions influence climate events? Did the CI eruption influence the decline of the Neanderthals? Do we recognize significant changes in the abundance and diversity of faunal and floral species after exceptionally large eruptions?
- Establishment of an in situ monitoring systems. This investigation regards the providing of optimal conditions for the quasi-real-time evaluation of critical parameters to be used as proxies to define the caldera dynamics.

4 Rationale of the Campi Flegrei drilling proposal

The outcomes of the workshop provided a conceptual framework for a full proposal for the drilling of the CFc to be submitted to the IODP and ICDP (Fig. 4) as an Amphibious Drilling Proposal. The Campi Flegrei ADP will combine complementary research topics into a general view on collapse-resurgent calderas located along continental margins. The partly submerged setting of the CFc provides a unique marine stratigraphic archive for a detailed reconstruction of the timing and kinematics of individual structures and components of the volcanic system, under different forcing factors (internal vs. external) during the past 10⁶ years (Fig. 5).

The drilling is important to reconstruct the subsurface 3-D stratigraphic architecture, identify faults and volcanic or volcano-tectonic features, and obtain information on the hydrothermal discharge areas and thermal structure. It also provides valuable information concerning the eruptive history of volcanoes and the dynamics of eruptions with different intensity. Previous research in active volcanic areas has shown that drilling can be fundamental in clarifying and constraining structural interpretations based on geophysical data alone. For instance, drilling at the Kakkonda geothermal field in Japan revealed the steep permeability and lithological gradients where magmatic and hydrothermal regimes interact (Saito et al., 1997; Nakada, 2013). The IDDP-1 (Iceland



Figure 4. Flow chart illustrating the history of site-surveying activity and nurturing of the continental (ICDP) and marine (IODP) components of the Campi Flegrei Amphibious Drilling Proposal.

Timeseelee	Forcing		
Timescales	Internal	External Sea-level change Tectonics	
Active – historical - Unrest, restlessness - Eruptions	Fluids Hydrothermalism Magmatism (shallow / deep)		
10 ³ – 10 ⁴ yrs - Post-caldera recovery, resurgence - Uplift, eruptions	Magmatism (shallow)	Sea-level change Tectonics Crustal detumescence	
10 ⁵ – 10 ⁶ yrs - Volcanic-magmatic history - Cyclicity, periodicity	Magmatism (flare-up) Crust / mantle	(Sea-level change?) Tectonics	

Figure 5. Timescales and forcing involved in the Campi Flegrei Amphibious Drilling Proposal (ADP).

Deep Drilling Project) well at Krafla in Iceland intercepted rhyolite melt in the region where geophysics had implied it would be absent (Elders et al., 2011). In the Long Valley Exploratory Well, a maximum temperature of ~ 100 °C was measured where the presence of shallow magma or at least very hot rock was assumed prior to drilling. Moreover, the use of selected drill holes as observatories provides an additional advantage for the in situ monitoring of volcanic and/or hydrothermal activity.

Drilling into active volcanic areas is not completely without risk, even though previous experiments in Japan and Iceland (Saito et al., 1997; Nakada, 2013; Elders et al., 2011) suggest that the chance of triggering a volcanic eruption is extremely small. The main risks are the hazards of possible underground pressure blowouts, meeting zones of fluid loss and material failure, and, more rarely, surface emission of liquefied sediments and steam (e.g., Sawolo et al., 2009). Documented experience indicates, however, that these risks can be significantly reduced or prevented by applying the appropriate mitigation techniques (e.g., blowout preventer systems) (e.g., Homuth et al., 2010).

4.1 Drilling at Campi Flegrei

The workshop was successful in identifying a number of relevant topics and questions, whose response may solve fundamental problems related to the caldera volcanism.

- a. The Campi Flegrei caldera represents an ideal example of an active caldera located in a shallow-water setting (< 200 m water depth). Other ODP Legs (Ocean Drilling Program) (e.g., ODP Leg 157: Gran Canaria and Madeira Abyssal Plain) and IODP Expeditions (e.g., IODP Expedition 340: Lesser Antilles Volcanism and Landslides) have focused on volcanic islands in deep oceanic settings. Campi Flegrei provides a unique opportunity to obtain a high-resolution stratigraphic archive of explosive, effusive, and extrusive volcanism, volcano-tectonic dynamics, and unrest. Moreover, the proximal marine setting of the CFc documents the primary deposition and reworking of pyroclastic currents and fall deposits as components of the continental shelf slope system.
- b. Campi Flegrei is a primary site to unravel the timing, structure, and evolution of caldera resurgence and unrest based on the geological record of marine strata. The mixed marine siliciclastic–volcaniclastic depositional architecture of the caldera fill provides a unique opportunity to document the pattern, timing, and rates of deformation related to resurgence since the Late Pleistocene. The last two millennia of documented unrest also provide further constraints in reconstructing time series of deformation onshore and offshore.
- c. The Campi Flegrei caldera generated the largest explosive eruptions in Europe during the Late Quaternary. The results of drilling and well logging will have a high potential impact on paleoenvironmental–paleoclimatic reconstruction. The coupling of proximal drill sites off Pozzuoli Bay with the results from the drilling of the distal stratigraphic record will also help in reconstructing the dispersal and erosive patterns of co-ignimbritic tephras. Also, this record could provide some insights into the puzzling issue of the apparent causal relationships between the environmental effects of the CI eruption and the final decline of Neanderthals.
- d. Drilling off the shore of the Campi Flegrei caldera will help investigate the interaction between the magmatic and hydrothermal systems and the occurrence of a wide range of subaerial-to-submarine features from

monogenic volcanoes to hydrothermal vents. The apparent difference between the onshore and offshore evolution may be related to changing magma–water interactions under saturated conditions within the mixing zone between the phreatic and marine pore waters, and this can only be investigated in detail by an onshore– offshore drilling transect. Long-term borehole monitoring of physical, chemical, and microbiological parameters may additionally provide a chance to identify the potential precursors to eruptions for the purpose of risk mitigation.

4.2 Amphibious drilling

The half-submerged setting of the CFc provides an opportunity to integrate results from offshore and onshore drillings and available marine geology and volcanological data. A deep, onshore borehole ($\sim 3 \text{ km}$) will allow the processes responsible for the recent unrest to be investigated at depth through the determination of rock physical properties, magma-water interaction, and water fluid chemical and physical exchanges. For instance, (1) extrapolated temperature measurements can be used to detect the depth of magmatic intrusions and the hydrothermal system; (2) the in situ chemical composition of fluids will provide information on rapid changes in the magmatic-hydrothermal system; and (3) deep monitoring systems will be deployed and incorporated in the existing network of the INGV-Napoli (Istituto Nazionale di Geofisica e Vulcanologia) (Osservatorio Vesuviano) to enable real-time tracking of such changes.

At the same time, a robust site-survey database, consisting of several multi-frequency (even 3-D) reflection seismic datasets (both single-channel and multichannel) has been acquired since 2008, yielding high-resolution images of the uppermost 100 m of the crust as well as new images to a depth of 1-2 km. Such a comprehensive database will enable a precise selection of offshore drill sites and guidance regarding deviated onshore drilling. The recent discovery of previously unknown volcanic structures and hydrothermal vents offshore offers a high potential for an integrated stratigraphic reconstruction (e.g., Sacchi et al., 2009, 2014; Passaro et al., 2016; Steinmann et al., 2016, 2018). The combined observations and data call for a joint offshore and onshore drilling program in order to (a) obtain an improved chronostratigraphic correlation between intra-caldera and extra-caldera products and (b) understand the origin of caldera collapse and the mechanisms of resurgence. An ideal drilling strategy at Campi Flegrei would therefore include the following onshore and offshore coordinated components:

a. Onshore drilling and well logging down to a depth in the order of 3000 m to investigate the caldera deep structure and associated deep magmatic-hydrothermal system. The drilling will provide the opportunity to investigate the processes responsible for the recent unrest episodes

at depth, thereby allowing for a reliable evaluation of the hazard. This component includes (1) the acquisition of physical-chemical parameters of the geothermal system over the entire depth; (2) stress measurements (size and direction) at depth; (3) the permeability measurements at depth; (4) the extrapolation of temperatures in the supercritical layer to detect the depth of the magmatic temperature and locate magmatic intrusions; and (5) the determination of the physical, mechanical, and rheological parameters of deep rocks.

b. Offshore drilling down to a maximum depth of $\sim 1000 \text{ m}$ to investigate the shallow structural levels of the caldera fill and resurgence as well as to unravel the mechanisms of magma-water interaction as a function of depth. This component provides the opportunity to study an undisturbed sedimentary archive without the challenges posed on land by intense subaerial erosion or urbanization (i.e., inaccessibility). This implies a much higher potential for structural, geochronological, petrographic, and geochemical reconstruction. Hence, marine drilling will provide a complete high-resolution stratigraphic record which will (1) improve the chronostratigraphy of volcanic and sedimentation events and unlock the timing and structural style of the deformation associated with inner-caldera resurgence, (2) understand the climatic effects and the environmental impact of ignimbrite eruptions on life and ecosystems, and (3) investigate the impact of magmatic-hydrothermal processes with respect to hydrothermal vents and shallow degassing structures as well as submarine monogenetic volcanoes and intrusions.

5 Drilling objectives and borehole logging and monitoring strategies

The workshop participants suggested that the IODP component of the ADP proposal should address the integrated stratigraphy of the caldera fill and resurgence, petrology, fluid geochemistry, and architecture of shallow structural levels (< 1000 m depth), whereas the ICDP component should focus on rock-fluid properties, physical-chemical processes, and the geothermal system at greater depth (< 3000 m). The proposed drilling strategy includes one major onshore drilling, complemented by an amphibious drilling transect extending from the Campi Flegrei shoreline towards the SE border of Naples Bay, together with distal drill sites in the Adriatic and Ionian Seas (Figs. 2, 6-12 and Table 1). Down-hole logging and long-term borehole monitoring at selected drill sites of primary physical and chemical parameters, along with microbiological analysis of rocks and fluids, within a depth range with a maximum of 0.5–1.0 km, have been also included in the planned operations.

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Figure 6. Location of onshore and offshore drill sites included in the first-draft plan of the Campi Flegrei Caldera Amphibious Drilling Proposal discussed during the MagellanPlus workshop.



Campi Flegrei collapse-resurgent caldera

Figure 7. Illustrated section of the Campi Flegrei caldera structure indicating the targets of the proposed onshore (ICDP) drill site (CFDDP-01). The reconstruction is mostly based on geophysical and geological data, affected by large uncertainties. The depth limit critical water temperature is constrained by previous drillings in the area (AGIP, 1987). Modified after De Natale and Troise (2011).

Table 1. Summary of proposed onshore and o	ffshore drill sites for the Campi Flegrei	Caldera Amphibious Drilling	g Proposal (ADP).
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ADP component	Proposed drill site	Structural sector	Drilling targets	Drilling depth (m)	Remarks
ICDP	CFDDP-01	Caldera margin	Stratigraphically reconstruct and well log through the hydrother- mal system down to the brittle and ductile zone	~ 3000	Deviated well, directed from the eastern border of the caldera to- wards the caldera center at depth
IODP	CF-01	Caldera center	Sample the stratigraphic succes- sion of NYT caldera fill and pene- trate the structural caldera floor	~ 900	Deep offshore well within the caldera collapse area; maximum drilling depth to be agreed on with safety panels
IODP	CF-02	Flanks of the resurgent dome	Drill the post-15 ka caldera fill to reconstruct the timing of deforma- tion and uplift of the caldera resur- gence	~ 50	Unique place to study the timing of the deformation of a caldera resur- gent structure
IODP	CF-03	Caldera collar	Drill the subsurface magmatic in- trusion and hydrothermal vent off the shore of Bagnoli (12–8 ka)	~ 120	Subsurficial intrusion (6–4 ka); magma-water interaction; implica- tions for volcanic hazard
IODP	CF-04	Caldera periph- ery	Drill the CI deposits in the shal- low subseafloor of Procida Chan- nel	~ 100	Proximal facies of the CI
IODP	CF-05	Caldera periph- ery	Drill the stratigraphic succes- sion of the peri-caldera mono- genic volcano of Miseno Bank (> 120 ka)	~ 80	Pre-caldera volcanism
IODP	CF-06	Caldera periph- ery	Drill the stratigraphic succession of the peri-caldera volcanic ap- paratus of Penta Palummo Bank (> 120 ka)	~ 80	Pre-caldera volcanism
IODP	CF-07	Caldera struc- tural border	Drill the stratigraphic succession of Nisida Bank (ca. 18–15 ka)	~ 100	Caldera-related volcanism
IODP	CF-08	Caldera periph- ery	Drill the subsurficial intrusion and hydrothermal vent of Mt. Dolce– Pampano Bank (8–4 ka)	~ 250	Subsurficial intrusion (18–15 ka); magma-water interaction; implica- tions for volcanic hazard
IODP	CF-09	Caldera exter- nal periphery	Drill the volcaniclastic diapirs and mounds associated with hy- drothermal venting at Montagna Bank (15–5 ka)	~ 150	Soft-sediment deformation and vol- caniclastic diapirism triggered by overpressured fluids
IODP	CF-010 CF-011 CF-012 CF-013	Proximal caldera exterior	Drill the Upper Quaternary mixed siliciclastic–volcaniclastic succession of Naples Bay for a stratigraphic purpose	$\sim 200 \\ \sim 200 \\ \sim 200 \\ \sim 200 \\ \sim 200$	Drilling transect for the recovery of a composite stratigraphic sec- tion; proximal stratigraphic record of Campi Flegrei eruptions
IODP	CF-014	Distal caldera exterior	Sample distal products of major explosive events from Campi Fle- grei and other eruptive centers of the Campanian district	~ 100	Distal tephrostratigraphy

0 -	Lithostratigraphy - facies interpretation				
0		Whitish pumice and scoria, with lava and tuff fragments, locally altered with reddish colors - Subaerial pyroclastic deposits, including the products of Agnano M. Spina (< 5 ka)			
50 -					
100 -		Lava, tuff, pumice and scoria (angular and rounded) fragments with variable abundance of siliceous fossils (sponge spiculae, diatoms). Locally charcoal and wood fragments. Peat level at ca. 60 m - <i>Subaerial pyroclastic deposits, and</i> <i>subaqueous volcaniclastic</i> deposits (10-5 ka)			
150 -		Gray pumice and lithics - Pyroclastic deposits			
200 -	16.9 ka 大	Black scoria, glass shards, with subordinate pumice and green tuffs - <i>Pyroclastic sequence</i> <i>including Neapolitan Yellow Tuff (NYT) towards</i> <i>the top</i>			
200-	35.1 ka ★	Gray pumice with subordinate black scoria - Pyroclastic deposits			
300 -		Gray to greenish pumice and lapilli within a biotite and feldspar crystals-bearing ash matrix. Between 290 and 315 m the deposit includes abundant yellowish lapilli in lenses or levels - <i>Pyroclastic sequence including Campania</i>			
450 -	38.5 ka ★	- Ignimbrite (CI) towards the base			
400-	47.4 ka 米	Zeolitized gray tuff including pumice and scoria in coarse-to-fine ash matrix - Pyroclastic deposits (Gray Tuff)			

CFDDP Pilot borehole

Figure 8. Synthetic lithostratigraphy and facies interpretation of the succession sampled at the CFDDP pilot borehole (500 m) on the shore of Bagnoli in 2012 (modified after De Natale et al., 2016). Red asterisks indicate the depth and 40 Ar/ 39 Ar age of sampled K feldspars. See Figs. 2–3 and 6 for the location of the borehole site.

5.1 Proposed on-land drill site (ICDP component)

Caldera margin to center

By drilling a $\sim 3 \text{ km}$ long deviated well from the eastern caldera margin towards its center (site CFDDP-01), we will be able to obtain a reference stratigraphic succession of the CFc fill to the basement floor and conduct well logging

through the hydrothermal system down to the brittle and ductile zone (Figs. 2, 6–7 and Table 1).

Another important component of the on-land drilling will be the deployment of a network of in situ monitoring stations at depth to provide real-time insights into changes in the hydrothermal–volcanic system. Such information is crucial to understanding the ongoing unrest as well as to reliably assessing hazards and risks. The drilling of site CFDDP-01 will rely on the results of the 500 m deep pilot hole and associated well log data acquired by the INGV-Napoli in 2012 (Figs. 2, 6 and 8).

5.2 Proposed offshore drill sites (IODP component)

5.2.1 Caldera center – caldera fill, resurgent dome, and structural floor

The caldera fill represents a high-resolution archive of the post-caldera volcanic succession, as well as a record of the ground deformation caused by caldera resurgence. Hence, drilling the caldera fill will facilitate (1) the discovery of new insights regarding post-caldera volcanic history, (2) the reconstruction of the timing, duration, and conditions of caldera resurgence, and (3) understanding the caldera's subsurface structure. Penetrating the floor of the caldera (site CF-01) will provide conclusive information on the precaldera phase and caldera formation processes. Site CF-02 is designed to recover the stratigraphic succession that formed over the flanks of the resurgent structure, in order to provide ages and timing of volcano-tectonic deformation since the NYT caldera collapse (last 15 000 years) (Figs. 2, 6, 9–11 and Table 1).

5.2.2 Caldera collar - fractured, permeable zone

The annular depression ("collar") between the structural border of the NYT caldera and the inner-caldera resurgent dome is a highly fractured zone, characterized by ascending fluids and locally shallow magmatic intrusions. The area represents a remarkably permeable segment of the caldera structure and is a key location to study the interconnection between the deep magmatic–hydrothermal system and the surface and its role during caldera unrest. Site CF-03 is planned to drill through the shallow structural levels of the ring-fault zone of the NYT caldera collar down to small laccolith-like intrusion off the shore of Bagnoli (Figs. 2, 6, 9–10 and Table 1). Drilling operations will be limited to shallow depths (< 150 m) and will be only realized after stepwise monitoring of temperature and fluid pressure.

5.2.3 Caldera structural border – pre-NYT caldera vents and intrusions

The outer border of the CFc is characterized by a number of offshore vents, shallow magmatic intrusions, and ignimbrite



NYT Caldera - Campi Flegrei

Figure 9. Schematic reconstruction of the shallow structure (< 1 km) of the collapse-resurgent caldera associated with the eruption of the Neapolitan Yellow Tuff (NYT) and the location of the proposed offshore (IODP) drill sites CF-01 and CF-03.



Figure 10. Interpreted high-resolution multichannel seismic profile across the CF caldera center and proposed location of drill sites CF-01 and CF-03. M1–M4 are the inner-caldera marine siliciclastic units; CI is Campanian Ignimbrite; NYT is Neapolitan Yellow Tuff.

deposits ranging in age from ~ 120 to ~ 18 ka. These provide a spectrum of volcanic features produced by significant magma–water interaction. They include most of the volcanic banks of the southern periphery of Pozzuoli Bay and CI ignimbrite deposits occurring at shallow depths beneath

the seafloor, mostly between Procida and the mainland. Drill sites will aim to characterize the nature of these volcanic centers and units and their role in the onset of pre-CFc volcanism and the overall fluid circulation (as, for instance, Nisida Bank represents an active fluid vent). Drilling at site CF-04



Figure 11. High-resolution single-channel seismic profile across the southern slope of the NYT resurgent dome and proposed location of drill site CF-02. Unconformities labeled as uplift 1–5 are interpreted as the result of a series of seafloor deformation phases associated with phases of the deformation of the resurgent structure. Correlation of tephra layers sampled by gravity core C32 is after Sacchi et al. (2014).

(Procida Channel) will focus on the recovery of a proximal succession of the CI. Drill sites CF-05 and CF-08 have been proposed to recover stratigraphic successions from a series of volcanic banks, namely at sites CF-04 (Miseno Bank), CF-05 (Penta Palummo Bank), CF-06, (Nisida Bank), and CF-07 (Mt. Dolce–Pampano Bank) (Figs. 2, 6 and Table 1).

5.2.4 Proximal extra-caldera area - Bay of Naples

Significant fluid venting in the Bay of Naples is not restricted to the ring-fault zone of the CFc, but it also occurs outside the structural border of the caldera itself. This is the case of Montagna Bank, a sub-circular seafloor region SE of Pozzuoli Bay that was formed by the dragging and rising up of volcaniclastic diapirs (consisting mostly of unconsolidated pumice), due to pore fluid overpressure at depth and associated fluid migration towards the seafloor (Passaro et al., 2016). Site CF-08 has been designed to drill through the volcaniclastic diapirs of Montagna Bank to the roots of the unconsolidated sediments involved in this process. The Naples Bay half-graben also represents an expanded, undisturbed sedimentary succession with interbedded massive ignimbrite deposits (NYT and CI). An offset drilling (CF-10 to CF-13) is an opportunity to cover a large time span of 1 million years (exceeding the entire time span of volcanic activity) with a transect of relatively shallow ($\sim 200 \text{ m}$) drillings (Figs. 2, 6 and Table 1). This will provide novel insights into the overall eruption history of the entire Campi Flegrei area and its tectonic evolution. Also, by drilling large ignimbrite units from top to bottom (i.e., contact zone of ignimbrite and siliciclastic

units), their environmental impact and subsequent the recovery of life after major eruptions can be investigated.

5.2.5 Distal extra-caldera area – Ionian Sea

A distal drill site (CF-14) has been proposed to recover an undisturbed pyroclastic fallout archive allowing for an integrated tephrostratigraphic analysis of the entire eruptive history of the Campi Flegrei area and other eruptive centers in the Campanian region (Figs. 6 and 12).

5.3 Down-hole logging and borehole monitoring strategies

The well logging plan incorporates a wide spectrum of downhole measurements which are designed to acquire maximum in situ information on petrophysical and geomechanical properties, as well as enhance monitoring of the strain-stress conditions, active seismicity, and the hydrothermal system at depth. The main parameters to be measured include (1) natural gamma rays, radioactivity, and spectrometry; (2) resistivity; (3) spontaneous potential redox; (4) sonic log; (5) magnetic susceptibility; (6) hole diameter (caliper); (7) temperature; (8) oriented microresistivity; and (9) acoustic and ultrasonic borehole images. The use of long-term borehole observatories (e.g., down-hole broadband seismic stations equipped with newly developed opto-electronic strain sensors and advanced monitoring systems that incorporate multiple seals allowing zoned measurements of in situ physical, chemical, and biological properties) may be considered for sites CFDDP-01 and CF-01.



Figure 12. Proposed site CF-14 (Ionian Sea) for distal tephrostratigraphic reconstruction of Late Quaternary volcanism of the Campanian region (Crocitti et al., 2018; Di Donato et al., 2019). (a) Location map with an indication of the main DSDP (Deep Sea Drilling Project), ODP, and IODP archives (yellow for marine and green for terrestrial) of the central Mediterranean region; (b) dispersal maps of plinian and sub-plinian events (Bronk-Ramsey et al., 2015); (c) dispersal maps of moderately explosive eruptions (Sulpizio et al., 2014; Crocitti et al., 2018). Please note that the years given in AD in Fig. 12c correspond to those same years in the CE notation system.

Earth.

6 Concluding remarks

Every eruption is preceded by unrest, but not every unrest culminates in an eruption (Acocella et al., 2015; Newhall and Dzurisin, 1988). Understanding the driving forces of volcanic unrest and the role of magmatic–hydrothermal processes is thus crucial for reliable hazard assessment. During the MagellanPlus workshop, all participants agreed that the CFc represents an ideal natural laboratory to study the interaction among volcanic, hydrothermal, marine, and volcanotectonic processes. The amphibious Campi Flegrei drilling project, involving a deep ICDP and shallower IODP drillings, will address fundamental aspects including phreatic and hydromagmatic volcanism, caldera formation and subsequent structural resurgence and post-caldera volcanism, fallout and ignimbrite stratigraphy, hydrothermal-magma interactions, mechanisms of volcanic unrest, and volcano-tectonic cou-

ermal proent. During ed that the giuseppe.denatale@ingv.it); offshore data are available at the Fac-

giuseppe.denatale@ingv.it); offshore data are available at the Faculty of Geosciences of the University of Bremen (Volkhard Spiess: vspiess@uni-bremen.de) and CNR-ISMAR, Naples (Marco Sacchi: marco.sacchi@cnr.it).

pling. The results will significantly advance our understanding of the most complex forms of volcanic structures on

Author contributions. MSa, GDN, VS, and LS jointly organized the workshop. MSa, GDN, LS, CK, and SDS drafted the paper. MSa, LS, and DI created the figures. All co-authors jointly contributed to the formulation of the concepts, scientific questions, and drilling/logging strategies discussed in the paper, according to their expertise: volcanology (SDS, NG, CS, HUS, MSu, and GV), petrology (LF), physical volcanology and volcanic hazards (GDN, VA, CK, AF, SP, RS, and CT), integrated stratigraphy (MSa, FM, and MV), structural geology of volcanic margins (VA, GV, and FP), tephrochronology (DI, PP, and ST), marine geophysics (VS, LS, SP, FP, and MC), and borehole logging (MJJ).

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. This report summarizes the results of the MagellanPlus workshop "Structure and Evolution of Magmatic and Hydrothermal Systems in offshore collapse-resurgent calderas – Development of an IODP Drilling Proposal at Campi Flegrei (Eastern Tyrrhenian Margin) linking to active ICDP Drilling Initiatives" (25–28 February, Naples, Italy). The authors wish to thank the two anonymous reviewers for their critical comments and suggestions on the early version of the manuscript.

Financial support. This research has been supported by the European Consortium for Ocean Research Drilling (ECORD) through the ECORD/ICDP MagellanPlus Workshop Series Programme (grant no. 2017.2376/4/AF).

Review statement. This paper was edited by Tomoaki Morishita and reviewed by two anonymous referees.

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