

## Origin of methane and light hydrocarbons in the gas manifestations of Greece

Daskalopoulou K.<sup>1-2\*</sup>, Calabrese S.<sup>1</sup>, Fiebig J.<sup>3</sup>, Grassa F.<sup>4</sup>, Kyriakopoulos K.<sup>2</sup>, Longo M.<sup>4</sup>, Parello F.<sup>1</sup>, Tassi F.<sup>5</sup>, D'Alessandro W.<sup>4</sup>

<sup>1</sup> Università degli Studi di Palermo, DiSTeM, via Archirafi, 36, 90123, Palermo, Italy, kyriaki.daskalopoulou@unipa.it,

<sup>2</sup> National and Kapodistrian University of Athens, Dept. of Geology and Geoenvironment, Panepistimioupolis, Ano Ilissia, 15784, Athens, Greece

<sup>3</sup> Goethe-Universität, Institut für Geowissenschaften, Altenhöferallee 1, 60438 Frankfurt am Main, Germany

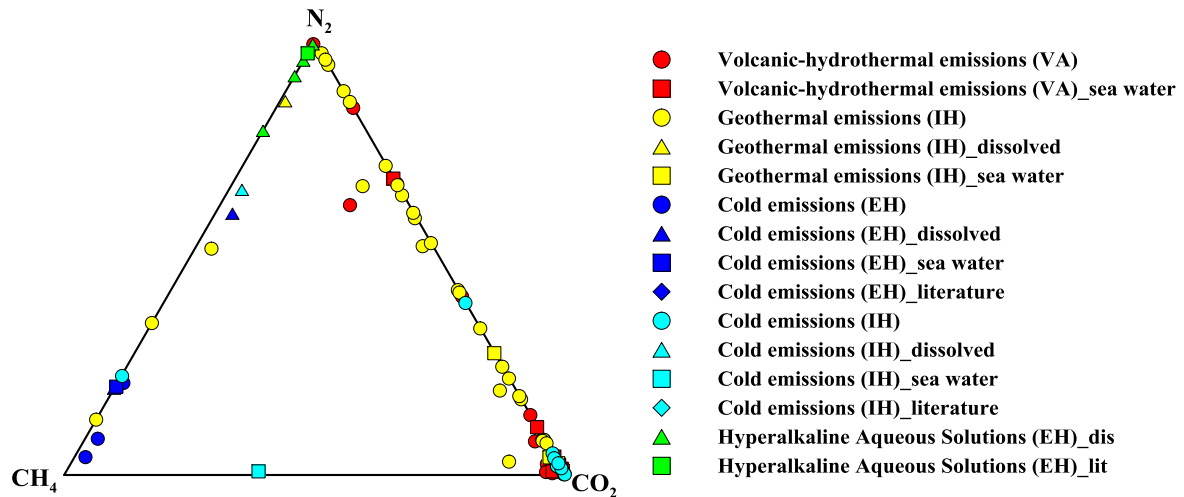
<sup>4</sup> Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Palermo, via Ugo la Malfa 153, 90146, Palermo, Italy

<sup>5</sup> Università degli Studi di Firenze, Dipartimento della Terra, via G. La Pira 4, 50121, Florence, Italy

The geologic emissions of greenhouse gases (CO<sub>2</sub> and CH<sub>4</sub>) give an important natural contribution to the global carbon budget. However, the contribution of these emissions to the global carbon cycle and their possible role on the climate change remain still poorly quantified (Guliyev and Feizullayev, 1997; Milkov, 2000; Etiope et al., 2015 and references therein). Methane, the most abundant organic compound in Earth's atmosphere, may be created either from existing organic matter or synthesized from inorganic molecules. Accordingly, it can be differentiated in two main classes: a) biotic (either microbial or thermogenic) and b) abiotic.

For this study, 115 gas samples of fumarolic, thermal and cold discharges from all over the Hellenic territory were collected and both chemical (CO<sub>2</sub>, H<sub>2</sub>S, CH<sub>4</sub>, N<sub>2</sub>, O<sub>2</sub>, Ar, H<sub>2</sub> and light hydrocarbons) and isotopic ( $\delta^{13}\text{C-CO}_2$ ,  $\delta^{13}\text{C-CH}_4$ ,  $\delta\text{D-CH}_4$ ) analyses were performed, in order to investigate the genetic processes that produced CH<sub>4</sub> in fluids related with the complex geodynamic setting of Greece. On the basis of the spatial distribution of the gas discharges and their type of emission, the whole dataset was subdivided into 3 main “domains”, as follows: 1) Volcanic Arc (VA) - 34 samples; 2) External Hellenides (EH) - 23 samples of cold emissions and of hyperalkaline aqueous solutions; 3) Internal Hellenides (IH) - 62 samples of cold and geothermal emissions. Almost each group is characterized, as long as subdivided in 3 groups based on the type of emission (on-land free or dissolved gases and subaqueous gases) and a 4<sup>th</sup> group includes literature data.

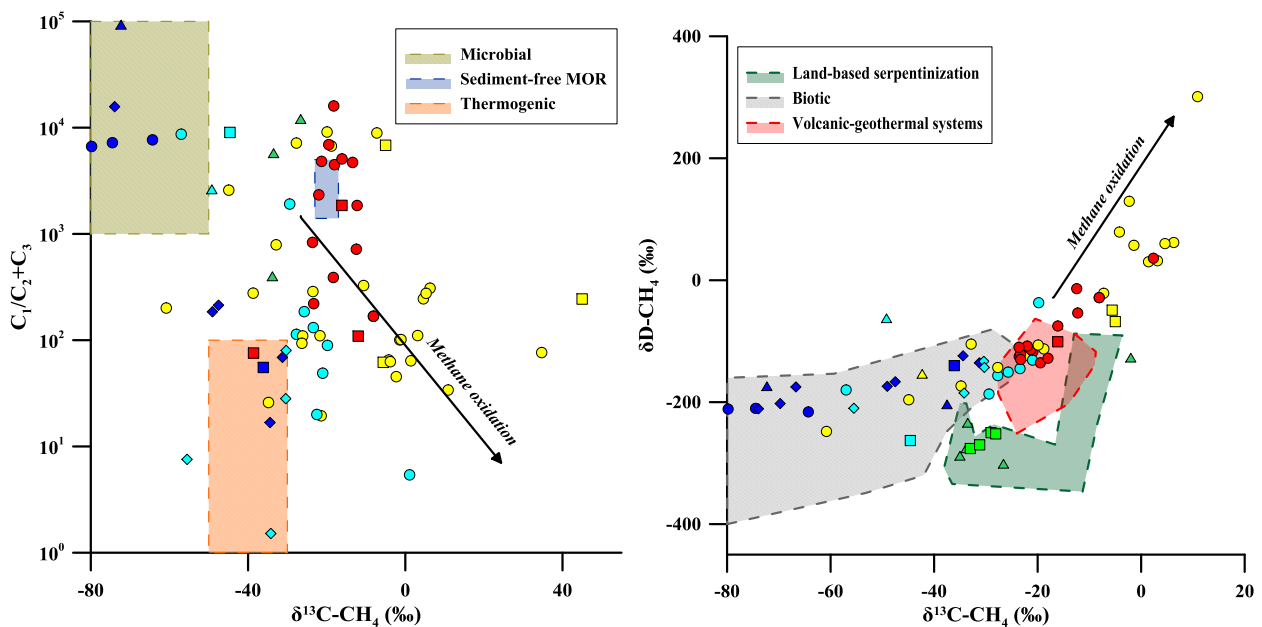
Samples collected from cold manifestations in the internal Hellenides and the Volcanic Arc mostly have CO<sub>2</sub> as the main species, whereas gases associated with hyperalkaline waters are N<sub>2</sub> dominated sometimes with significant CH<sub>4</sub> contributions. In the cold manifestations collected from the external Hellenides the prevailing gas is CH<sub>4</sub>. The remaining gas samples collected from the geothermal emissions that occur in the Hellenic territory are showing a mixed N<sub>2</sub> - CO<sub>2</sub> composition (Figure 1).



**Figure 1:** Chemical composition of the gas manifestations in Greece. VA = volcanic arc; IH = internal Hellenides; EH = external Hellenides

The present study highlights a widespread continental and underwater degassing activity along the Hellenic territory. Both chemical and isotopic compositions of CH<sub>4</sub> underscore the different primary sources and the secondary post-generic processes (oxidation) that can significantly affect the origin of this gas compound (Figure 2a, 2b). Hydrocarbons in the CH<sub>4</sub>-dominated gases from the external Hellenides are showing a clear biotic origin. In particular, those collected in the Gavrovo-Tripolis zone are showing a dominating thermogenic origin, whereas it is also noticeable that some of the samples of the Ionian zone are produced by both microbial activity and thermal maturation of sedimentary organic matter. The CO<sub>2</sub>-dominated gas discharges from the main geothermal systems of the Internal Hellenides and from the Volcanic Arc most likely predominantly contain abiogenic CH<sub>4</sub> deriving from CO<sub>2</sub> reduction. This process seems to have a lower effectiveness in producing higher hydrocarbons, such as C<sub>2</sub>H<sub>6</sub>, C<sub>3</sub>H<sub>8</sub> and C<sub>6</sub>H<sub>6</sub>. However, some of the CO<sub>2</sub>-rich gas discharges of the geothermal and volcanic-hydrothermal systems located in the sedimentary Pelagonian and the Gavrovo-Tripolis zones, seem to exhibit significant contributions from thermogenic sources. This is likely due to 1) low temperatures of the fluid reservoirs, i.e. too low to promote an efficient CO<sub>2</sub> to CH<sub>4</sub> conversion, and 2) the large amounts of organic matter that

were buried in the sedimentary formations and that is available for thermal degradation processes. The presence of abiogenic methane was also recognized in the hyperalkaline aqueous solutions that are issuing from the ophiolites of Othrys and Argolida (Etioppe et al., 2013; D'Alessandro et al., 2017). Most of the geothermal gases of central Greece (internal Hellenides) and some of the volcanic-hydrothermal ones show lower Bernard ratios ( $\text{CH}_4/[\text{C}_2\text{H}_6+\text{C}_3\text{H}_8]$  – Fig. 2a) and strongly positive isotopic ratios of  $\text{CH}_4$  ( $\delta^{13}\text{C}$  up to +45‰ and  $\delta\text{D}$  up to 301‰ – Fig. 2). Such chemical and isotopic features can be explained by microbial oxidation of  $\text{CH}_4$ . In these environments microbes obtain energy from aerobic or anaerobic  $\text{CH}_4$  oxidation (Murrell and Jetten, 2009), preferentially consuming  $\text{CH}_4$  with respect to higher hydrocarbons and preferring light isotopes. This results in a noticeable decrease in the Bernard ratio and a progressive enrichment in heavy isotopes (both  $^{13}\text{C}$  and D) in the residual  $\text{CH}_4$ .



**Figure 2:** a) Bernard classification diagram (Bernard et al., 1978) that is correlating the  $\text{CH}_4/(\text{C}_2\text{H}_6+\text{C}_3\text{H}_8)$  concentration ratios with the  $\delta^{13}\text{C}-\text{CH}_4$  ratios for the Hellenic gas discharges. Values for gases of biogenic origin (microbial and thermogenic) and from unsedimented mid-oceanic ridges are reported (McCollom and Seewald, 2007, and references therein) for comparison; b) Schoell binary diagram (Etioppe and Schoell, 2014) which is correlating  $\delta\text{D}-\text{CH}_4$  with  $\delta^{13}\text{C}-\text{CH}_4$  ratio for the Hellenic gas discharges.

## References

BERNARD, B.B., BROOKS, J.M., SACKETT, W.M., 1978. A geochemical model for characterization of hydrocarbon gas sources in marine sediments. Offshore Technology Conference, Houston, USA, 435-438.

- D'ALESSANDRO, W, DASKALOPOULOU, K., CALABRESE, S., BELLOMO, S., 2017. Water chemistry and abiogenic methane content of a hyperalkaline spring related to serpentinization in the Argolida ophiolite (Ermioni, Greece). *Marine and Petroleum Geology*, doi: 10.1016/j.marpetgeo.2017.01.028.
- ETIOPE, G., 2015. *Natural Gas Seepage. The Earth's Hydrocarbon Degassing*. Springer International Publishing Switzerland, e- book. DOI 10.1007/978-3-319-14601-0.
- ETIOPE, G., SCHOELL, M., 2014. Abiotic gas: atypical but not rare. *Elements* 10, 291-296.
- ETIOPE, G., TSIKOURAS, B., KORDELLA, S., IFANDI, E., CHRISTODOULOU, D., PAPANTHEODOROU, G., 2013. Methane flux and origin in the Othrys ophiolite hyperalkaline springs, Greece *Chemical Geology* 347, 161–174.
- GULIYEV, I. S. AND FEIZULLAYEV, A. A., 1997. *All about mud volcanoes*. Baku Pub. House, NAFTA-Press, 120.
- MCCOLLOM, T.M., SEEWALD, J.S., 2007. Abiotic synthesis of organic compounds in deep- sea hydrothermal environments. *Chemical Reviews* 107, 382-401.
- MILKOV, A. V., 2000. Worldwide distribution of submarine mud volcanoes and associated gas hydrates, *Marine Geology* 167 (1 – 2), 29 – 42.
- MURRELL, C. J. AND JETTEN, M. S. M., 2009. The microbial methane cycle, *Environ. Microbiol. Reports* 1, 279–284.