

FULL PAPER

Open Access



# Total (fumarolic + diffuse soil) CO<sub>2</sub> output from Furnas volcano

M. Pedone<sup>1,2\*</sup>, F. Viveiros<sup>3</sup>, A. Aiuppa<sup>1,2</sup>, G. Giudice<sup>2</sup>, F. Grassa<sup>2</sup>, A. L. Gagliano<sup>2</sup>, V. Francofonte<sup>2</sup> and T. Ferreira<sup>3</sup>

## Abstract

Furnas volcano, in São Miguel island (Azores), being the surface expression of rising hydrothermal steam, is the site of intense carbon dioxide (CO<sub>2</sub>) release by diffuse degassing and fumaroles. While the diffusive CO<sub>2</sub> output has long (since the early 1990s) been characterized by soil CO<sub>2</sub> surveys, no information is presently available on the fumarolic CO<sub>2</sub> output. Here, we performed (in August 2014) a study in which soil CO<sub>2</sub> degassing survey was combined for the first time with the measurement of the fumarolic CO<sub>2</sub> flux. The results were achieved by using a GasFinder 2.0 tunable diode laser. Our measurements were performed in two degassing sites at Furnas volcano (Furnas Lake and Furnas Village), with the aim of quantifying the total (fumarolic + soil diffuse) CO<sub>2</sub> output. We show that, within the main degassing (fumarolic) areas, the soil CO<sub>2</sub> flux contribution (9.2 t day<sup>-1</sup>) represents a minor (~15 %) fraction of the total CO<sub>2</sub> output (59 t day<sup>-1</sup>), which is dominated by the fumaroles (~50 t day<sup>-1</sup>). The same fumaroles contribute to ~0.25 t day<sup>-1</sup> of H<sub>2</sub>S, based on a fumarole CO<sub>2</sub>/H<sub>2</sub>S ratio of 150 to 353 (measured with a portable Multi-GAS). However, we also find that the soil CO<sub>2</sub> contribution from a more distal wider degassing structure dominates the total Furnas volcano CO<sub>2</sub> budget, which we evaluate (summing up the CO<sub>2</sub> flux contributions for degassing soils, fumarolic emissions and springs) at ~1030 t day<sup>-1</sup>.

**Keywords:** Carbon dioxide flux, Fumaroles, Soil diffuse degassing, Furnas volcano

## Background

Volcano-hosted hydrothermal systems are the source of sizeable carbon dioxide (CO<sub>2</sub>) emissions, either vented by hydrothermal steam vents (Chiodini et al. 1998) or diffusively released by degassing soils (Chiodini et al. 1999; Rogie et al. 2001; Werner et al. 2008). While considerable work has been spent in the past to estimate the soil CO<sub>2</sub> flux from hydrothermal areas (e.g., Chiodini et al. 1999, 2001a; Hernández et al. 2001; Werner et al. 2008; Viveiros et al. 2010), far less is known on the CO<sub>2</sub> output from hydrothermal fumarolic vents (Werner et al. 2000; Fridriksson et al. 2006; Aiuppa et al. 2013; Pedone et al. 2014a, b), which are technically more difficult to study. Consequently, the total (fumarolic + soil degassing) budget remains unconstrained for most hydrothermal systems in our planet (with a few exceptions; Aiuppa et al. 2013).

The aim of the present work is to provide more information on the fumarolic vs. diffusive contribution to the hydrothermal CO<sub>2</sub> output. Our test site is Furnas volcano, a quiescent polygenetic volcano located in the eastern part of São Miguel island, in the Azores, an archipelago of nine volcanic islands located in the North Atlantic Ocean at the triple junction between American, Eurasian, and Nubian plates (Searle 1980). Furnas volcano has frequently been active in the Holocene (the oldest volcanic products are dated back 100,000 years BP; Moore 1990). The last “magmatic” eruption occurred in 1630 (Cole et al. 1995). In recent times, hydrothermal explosions have re-occurred (in 1840–1841, 1944, and 1990) from the hydrothermal vent (named “Asmodeu”) belonging to the Furnas Village fumarolic field (Ferreira, T: Contribuição para o estudo das emanações gasosas associadas a processos de vulcanismo no arquipélago dos Açores, unpublished Master thesis). Hydrothermal activity is widespread on the island and includes soil diffuse degassing areas (Ferreira et al. 2005; Viveiros et al. 2010), steaming ground, thermal springs, cold CO<sub>2</sub>-rich springs, and low-temperature fumaroles (95–100 °C), mostly concentrated

\* Correspondence: pedone\_maria@libero.it

<sup>1</sup>DiSTeM, Università di Palermo, via Archirafi, 36, Palermo 90123, Italy

<sup>2</sup>Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Palermo, via Ugo La Malfa, 153, Palermo 90146, Italy

Full list of author information is available at the end of the article

inside the Furnas caldera (where three main fumarolic fields are observed; Viveiros et al. 2010; Caliro et al. 2015). Further studies have been done since the early nineties to study CO<sub>2</sub> diffuse emissions. The first soil CO<sub>2</sub> surveys (Baubron et al. 1994; Baxter et al. 1999) in the Furnas caldera identified a CO<sub>2</sub> degassing area in the proximity of Furnas village. Recently, Viveiros et al. (2010, 2012) estimated the soil CO<sub>2</sub> fluxes emitted from the Furnas volcanic system using the accumulation chamber method (Chiodini et al. 1998); this led to identifying the presence of several diffuse degassing structures (DDS). The diffuse hydrothermal-volcanic CO<sub>2</sub> output from Furnas volcano (Furnas caldera and the southern Ribeira Quente village area) was estimated at ~968 t day<sup>-1</sup> (Viveiros et al. 2010), and the groundwater CO<sub>2</sub> transport was evaluated at ~12 t day<sup>-1</sup> (Cruz et al. 1999). As for the majority of the hydrothermal system worldwide, the fumarolic CO<sub>2</sub> output is unknown.

In this study, we use a tunable diode laser spectrometer (TDLS) to estimate, for the first time, the fumarolic output of volcanic/hydrothermal CO<sub>2</sub> at Furnas volcano. The TDLS technique is based on measuring the absorption of IR radiation (at specific wavelengths) by a target gas, and can suitably be adapted to measure the flux of volcanic CO<sub>2</sub> from low-temperature hydrothermal manifestations (Pedone et al. 2014a, b), where the use of traditional UV spectroscopy remote-sensing techniques is prevented by the absence of SO<sub>2</sub>. TDLS employs a light source of very narrow line-width that is tunable over a narrow wavelength range. In other words, tunable diode laser (TDL) steams on absorption spectroscopy using a single isolated absorption line of the target species, allowing positive identification and unambiguous measurement of complex gas mixtures. A major disadvantage is that TDLS applications are better suited to accurate measurement of a specific target gas (known to be present in the atmosphere) than for identification of previously unidentified species (Pedone et al. 2014a). In addition, the quality of the measurements can be limited in highly condensed, optically thick fumarolic plumes.

The fumarole observations were complemented by simultaneous soil CO<sub>2</sub> measurements with an accumulation chamber. This is the first time in which two methodologies are applied together to evaluate their relative CO<sub>2</sub> contributions to the total CO<sub>2</sub> output. Our results, while limited to only one single hydrothermal system, offer new information to understand the modes of hydrothermal carbon release. In addition, the sulfur output from the fumarolic system is quantified based on measurement of the fumarole CO<sub>2</sub>/H<sub>2</sub>S ratios (obtained with a Multi-GAS detector).

#### Site description

The two most important fumarolic fields on Furnas are referred to as “Furnas Lake” and “Furnas Village” (see

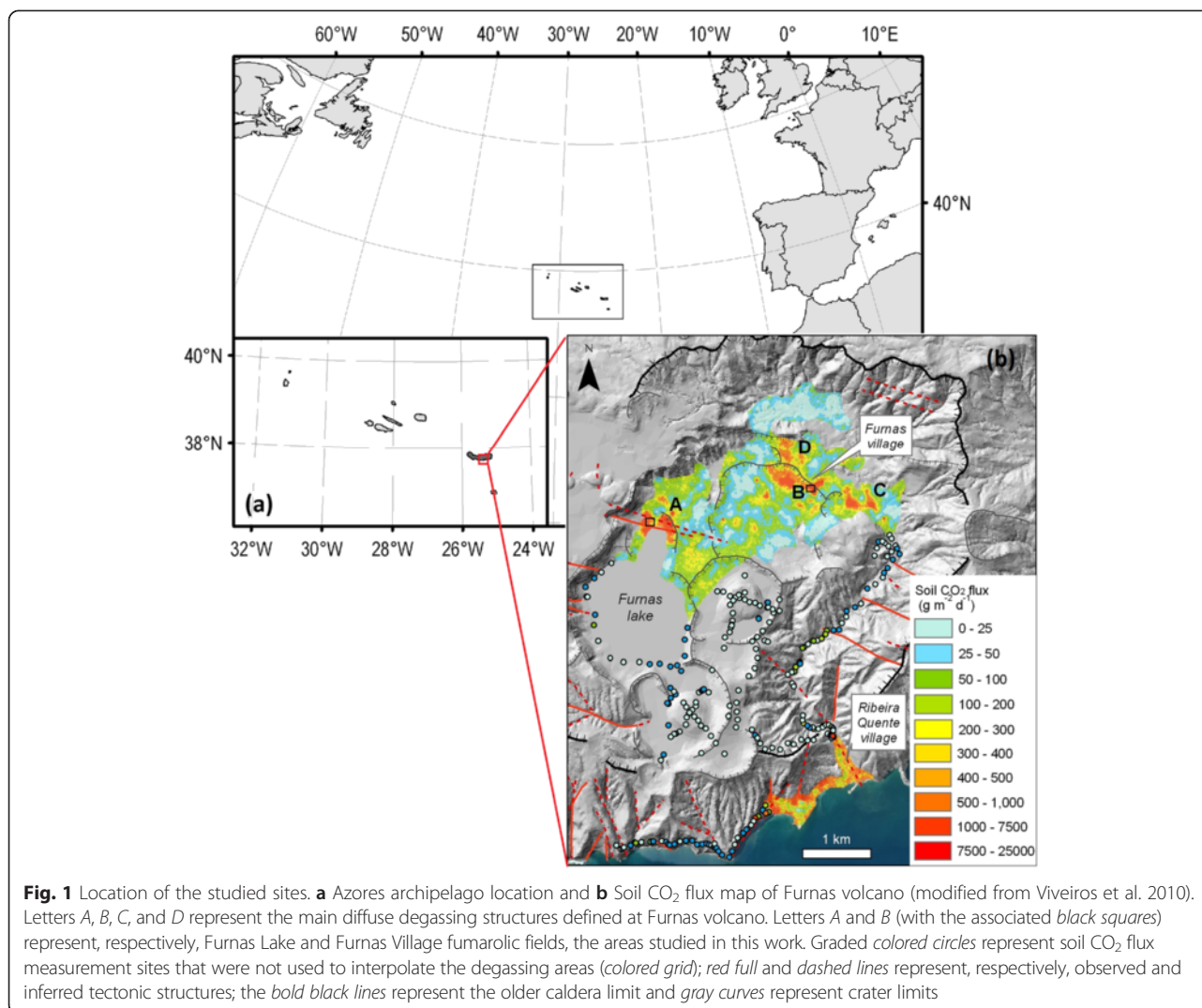
Fig. 1). A small steaming ground exists in the eastern part of the caldera, and the steam emissions are visible in the southern flank of the volcano (Ribeira Quente village).

Our field campaign was carried out in August 2014 in the two main fumarolic fields: Furnas Lake, on the north side of the lake, and Furnas Village, in the eastern side of the caldera (Fig. 1). Figure 2 shows a view of the Furnas Lake area during the campaign carried out on 19 August 2014. The TDL was used to measure the CO<sub>2</sub> concentration in air close to and/or above three degassing vents (Fig. 2): a vigorously degassing vent with boiling water jet (WJ), close to mirror position 3 in Fig. 2); two smaller fumaroles located in the northwestern part of the area (ST, close to mirror position 2 in Fig. 2); and a mud pool degassing vent (MP, close to mirror position 5 in Fig. 2). Furnas Village (Fig. 3) is the largest fumarolic field in the Furnas caldera. Several moderate to large fumaroles are active within an area of several hundred square meters. “Caldeira Grande” and “Caldeira do Asmodeu” are the main hydrothermal vent manifestations, and are referred to as “CG” and “CdA” in the subsequent text and figures. These fumarolic fields have hydrothermal origin (Ferreira and Oskarsson 1999), with outlet temperatures lower than 100 °C (~97 °C at Lake and ~99 °C at Village; Caliro et al. 2015). The discharge of these fumaroles derive from the shallow hydrothermal aquifers originating from heating of infiltrated local groundwater. The main components are water vapor, carbon dioxide, hydrogen sulfide, nitrogen, hydrogen, oxygen, methane, and argon (Ferreira and Oskarsson 1999; Ferreira et al. 2005; Caliro et al. 2015).

#### Methods

Tunable diode lasers are increasingly used in environmental monitoring applications (Gianfrani et al. 1997a) and for volcanic gas observations (Gianfrani et al. 1997b, 2000; De Natale et al. 1998; Richter et al. 2002). Pedone et al. (2014a, b) recently reported on the first direct observations of the volcanic CO<sub>2</sub> flux by using a portable tunable diode laser (TDL) system.

Like in previous studies (Pedone et al. 2014a, b), we used a GasFinder 2.0 Tunable Diode Laser (produced by Boreal Laser Inc.), a transmitter/receiver unit operating in the 1.3–1.7 μm wavelength range. GasFinder 2.0 is designed to measure CO<sub>2</sub> concentrations over linear open-paths of <1 km. In order to achieve this, radiation emitted by the IR laser transmitter propagates to a set of gold-plated retroreflector mirrors, where it is reflected back to the receiver and focused onto a photodiode detector. The CO<sub>2</sub> column amount (in ppm-m) along each optical path is calculated by the spectral analysis of reflected light, converted into electrical waveform, and processed by using the procedure described in Tulip (1997). CO<sub>2</sub> column amounts are converted into average



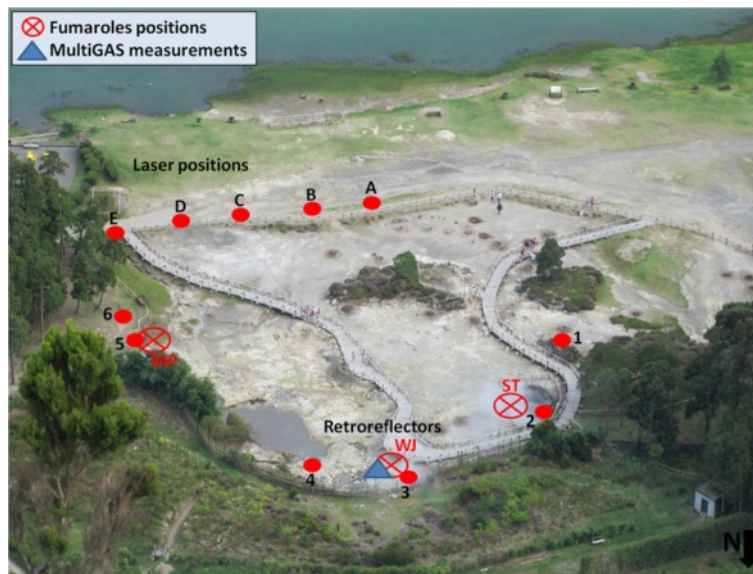
CO<sub>2</sub> concentrations (in ppm) along the path by considering path lengths (measured with an IR manual telemeter, 1 m resolution).

In the field, the GasFinder was set to measure CO<sub>2</sub> concentrations at a 1-Hz rate (Pedone et al. 2014a), along an optical (transmitter-retroreflector) path positioned at ~1.20 m height above ground level. Alignment between the laser unit and the retro-reflectors was optimized using a red visible aiming laser and a sighting scope. The GasFinder acquired for several consecutive hours at Furnas Lake (on 19 August 2014) and Furnas Village (on 22 August 2014), totally ~3600 and ~3200 successful readings, respectively. The plumes were fairly transparent during the observations, and the main plume dispersals directions are indicated by rose diagrams in Figs. 4 and 5. Background CO<sub>2</sub> concentrations were obtained in each of the measurement days/sites by pointing the laser beam toward a mirror, positioned upwind of the fumarolic area. These background values

were subtracted in the calculation of integrated column amount (ICA) (see below).

In both days, the position of the GasFinder laser unit was sequentially moved (e.g., from positions A to E in Figs. 2 and 3) to scan the fumaroles' atmospheric plumes from different angles and viewing directions. The positions of transmitter-receivers were limited by time-logistic constraints (e.g., morphology and accessibility of the degassing areas); given the geometry of the optical paths, we admit some heterogeneities in our CO<sub>2</sub> maps (Figs. 4 and 5) may result from inhomogeneous, incomplete coverage of the degassing areas. During operations, the GasFinder was left to acquire data along each single GasFinder-retroreflector path for ~3–5 min, before being rotated to measure along the successive path (an additional documentation file shows more details of acquisition-paths (see A1 in Additional file 1).

We considered the possibility of combining the available data to derive a two-dimensional model of CO<sub>2</sub>



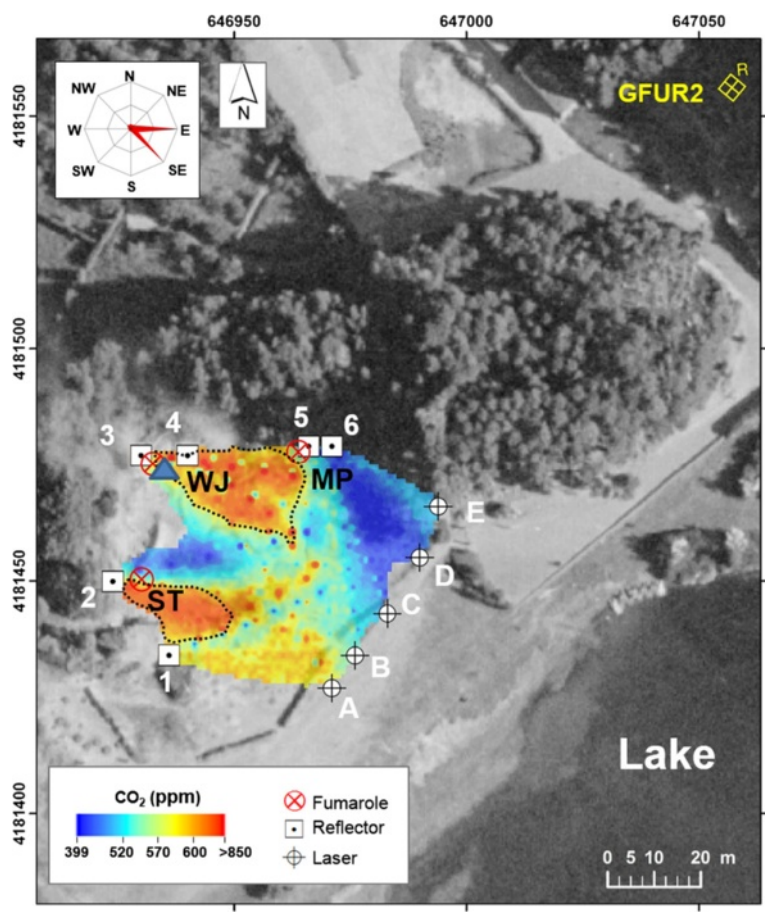
**Fig. 2** View of Furnas Lake fumarolic site from the top of northern part of the Furnas caldera. The positions of GasFinder unit and retro-reflectors are shown with letters and numbers, respectively. The position of Multi-GAS measurement-point (blue triangle) and the degassing vents "ST," "WJ," and "MP" (red crossed-circles) are also given

distribution, and ultimately—by combination with derived plume vertical transport speeds (Table 1)—to estimate the CO<sub>2</sub> output from the main investigated fumaroles. The initial step was to use a Matlab script (Pedone et al. 2014a) to obtain a 2D reconstruction of CO<sub>2</sub> concentrations in the investigated areas, starting from the raw GasFinder dataset. In order to begin calculation, the Matlab script was initialized with the coordinates

of GasFinder unit and retroreflector positions (Figs. 2 and 3). The additional input data was a column vector containing the mean CO<sub>2</sub> column amount (in ppm-m) obtained for each GasFinder-retroreflector path. With these inputs, the script performed a data inversion using a least-squares method, whose outputs were then processed with a stochastic simulation approach (see Additional file 1: Table S3 and S4), using GSLIB executables (Deutsch and



**Fig. 3** Furnas Village fumarolic site. The positions of GasFinder unit and retro-reflectors are shown with letters and numbers, respectively. The position of Multi-GAS measurement-point (blue triangle) and two main degassing vents "CG" and "CdA" (red crossed-circles) are also given. CG, Caldeira Grande, CdA Caldeira do Asmodeu



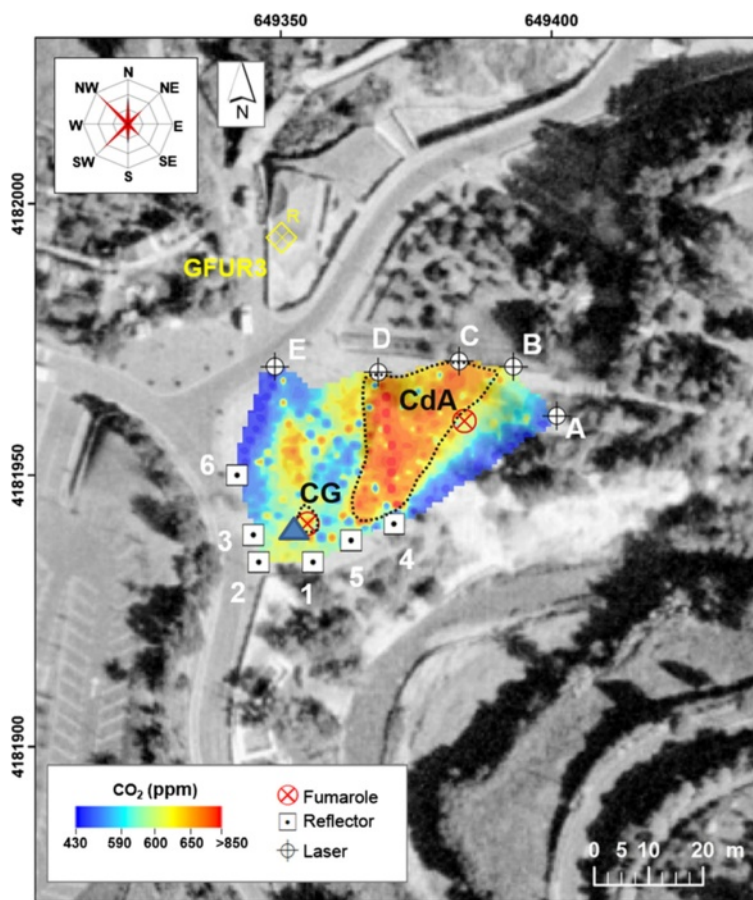
**Fig. 4** CO<sub>2</sub> concentrations (ppm) map at “Furnas Lake” survey (using TDL). GasFinder and retro-reflectors positions are shown with letters and numbers, respectively. Investigated fumaroles (WJ, ST, and MP; red crossed-circles), Multi-GAS measurement-point (blue triangle), and the GFUR2 meteo station positions are given. The principal directions of plume/wind dispersal are given (rose diagram). The CO<sub>2</sub> flux was estimated integrating the concentration values (>600 ppm) inside the black dashed lines, close to the main fumaroles and downwind dispersal direction

Journel 1998) to obtain the distribution maps of CO<sub>2</sub> concentrations, shown in Figs. 4 and 5. For each area, the dataset was initially converted by normal score transformation to follow a Gaussian distribution. This normal score-transformed dataset was then used to compute omnidirectional variograms, and finally interpolated with the sequential Gaussian simulations (sGs) method to produce 100 realizations. The maps of Figs. 4 and 5 were produced averaging results of the 100 simulations, using the E-type post-processes method (Cardellini et al. 2003). Grid resolution interpolation parameters are listed in Additional file 1: Table S5. Zonal statistics on the CO<sub>2</sub> distribution maps was obtained by using the ArcMap 9.3 (ESRI) Spatial Analyst tool (Additional file 1: Table S5).

To convert the CO<sub>2</sub> maps into a CO<sub>2</sub> flux, we performed 2D integration inside the areas covered by the fumarolic plumes—the boxes delimited by black dashed lines in Figs. 4 and 5. These integration areas were delimited based on visual (field) observations of plume transport direction, and were mapped so as to include

concentration data above fixed concentration thresholds (~600 ppm at lake and ~650 ppm at village). No integration was performed outside these areas, where the CO<sub>2</sub> contribution was due either to the degassing soils or smaller fumaroles (for which manifestations—as for mud pool “MP”—the gas speed is more difficult to assess). The so-obtained CO<sub>2</sub> ICAs were multiplied by vertical gas transport speed to obtain a CO<sub>2</sub> flux (Table 1). Plume vertical transport speed was estimated from recordings of a video camera, pointing toward the fumarolic vents, and acquiring sequences of images of the atmospheric plume at 25 frames per second. The sequences of frames were later post-processed to calculate the time-averaged transport speed of the plume, after converting camera pixels into distances (using a graduated pole, positioned close to the vent).

During the field campaigns, we used a Multi-GAS (Aiuppa et al. 2011, 2012 for information about setup and performance of the instrument) to measure the compositions of the main gas manifestations. In detail, the Multi-GAS was exposed to the atmospheric plumes



**Fig. 5** CO<sub>2</sub> concentrations (ppm) map at “Furnas Village” survey (using TDL). GasFinder and retro-reflectors positions are shown with letters and numbers, respectively. Investigated fumaroles (red crossed-circles) and Multi-GAS measurement-point (blue triangle), and the GFUR3 meteo station positions are given. The principal directions of plume/wind dispersal are given (rose diagram). The CO<sub>2</sub> flux was estimated integrating the concentration values (>650 ppm) inside the black dashed lines, close to the main fumaroles and downwind dispersal direction. CG Caldeira Grande, CdA Caldeira do Asmodeu

**Table 1** CO<sub>2</sub> fluxes (in t day<sup>-1</sup>) and standard deviation ( $\pm 1 \sigma$ ) calculated in the investigated areas based on TDL and accumulation chamber surveys. The plume transport vertical speed (in m s<sup>-1</sup>) is also given for each site. Fumarolic CO<sub>2</sub> outputs are given for each fumarole emissions (WJ and ST at Furnas Lake; CdA and CG at Furnas Village). Soil CO<sub>2</sub> flux values are given for the entire areas (Lake and Village, respectively). Total CO<sub>2</sub> emissions (fumarolic + soil) are also given

Site	Date	Gas speed	CO <sub>2</sub> flux <sup>a</sup>	CO <sub>2</sub> flux <sup>b</sup>	Total flux <sup>c</sup>
WJ	19-08	0.98 ± 0.07	17.6 ± 5.3	6.0 ± 0.2	35
ST	19-08	1.00 ± 0.02	11.4 ± 4.7	–	–
CdA	22-08	0.90 ± 0.18	17.8 ± 4.1	3.2 ± 0.2	24
CG	22-08	1.80 ± 0.19	3.00 ± 0.8	–	–

<sup>a</sup>Estimated from GasFinder datasets

<sup>b</sup>Estimated from accumulation chamber datasets

<sup>c</sup>Fumarolic + soil CO<sub>2</sub> emissions

of the main fumaroles (WJ at Furnas Lake and CG at Furnas Village; blue triangles in Figs. 2, 3, 4, and 5) to measure the concentrations of CO<sub>2</sub> (by NDIR spectroscopy), SO<sub>2</sub>, and H<sub>2</sub>S (by specific electrochemical sensors). The specific sensors mounted onboard the Multi-GAS were as follows: a Gascard Edinburgh Instruments infrared spectrometer for CO<sub>2</sub> (0-3000 ppmv range, with a resolution of 0.8 ppmv), a 0- to 200-ppmv SO<sub>2</sub> electrochemical sensor, and a 0- to 50-ppmv H<sub>2</sub>S electrochemical sensor (all from City Technology, and the manufacturer quoted a resolution of 0.5 ppmv). We also used a temperature-humidity Galltec sensor (*T* range, -30 to 70 °C; Rh range, 0–100 %). The Multi-GAS sensors were calibrated in the laboratory by using standard gas cylinders of concentrations within the sensor ranges (all in nitrogen matrixes). Laboratory tests indicate a typical measurement error in the CO<sub>2</sub>/H<sub>2</sub>S ratios of ≤20 %.

In both areas, we also performed a detailed soil CO<sub>2</sub> flux survey using a portable accumulation chamber (Chiodini et al. 1998). A chamber of a known volume is placed on the soil surface, the variation of CO<sub>2</sub> concentration is measured, and the CO<sub>2</sub> flux is calculated from the increase rate of the CO<sub>2</sub> concentration. The portable chamber is equipped with a LICOR LI-800 infrared CO<sub>2</sub> detector that measures CO<sub>2</sub> concentrations in the range 0–2 vol %. Calibration of the instrument was performed in the laboratory prior to field work. Considering that soil gas fluxes may be highly affected by meteorological conditions (e.g., Viveiros et al. 2008), measurements were performed during stable weather conditions and the CO<sub>2</sub> flux values recorded in a permanent soil CO<sub>2</sub> flux station (named GFUR2) used as control site; this station is installed close to Furnas Lake fumarolic field (Fig. 4). Soil CO<sub>2</sub> fluxes in the permanent station showed a coefficient of variation (ratio between standard deviation and average value) of 0.02 and 0.03 during the surveys carried out in 19 and 22 August 2014, respectively, at Furnas Lake and Furnas Village fumarolic fields. At Furnas Lake, a total of 124 measurements were carried out in an area of approximately 3666 m<sup>2</sup>. At Furnas Village, a total of 95 soil CO<sub>2</sub> flux measurements were carried out in an area of about 2300 m<sup>2</sup>. A sampling space varying between 3 and 5 m was used in these detailed surveys. The grid used was as regular as possible, considering the presence of steam vents, water streams, irregular topography, and some man-made structures in the study sites. The CO<sub>2</sub> flux values were interpolated using sGs (Deutsch and Journel 1998; Cardellini et al. 2003), which consists in the production of numerous simulations of the spatial distribution of the CO<sub>2</sub> flux. Considering that data has to follow normal distribution in order to apply sGs, normal score transformation was applied to the original data. Similarly as described above, in this case, omnidirectional variograms of normal scores for the two surveyed areas were also computed (Fig. 6). One-hundred realizations of the flux grid were created for each study site. Parameters used in the sGs procedure are available in A3 section of Additional file 1: Tables S6 and S7. The CO<sub>2</sub> output was calculated by

integrating the average values estimated from sGs over the area. The mean and the standard deviation computed for the 100 realizations are, thus, assumed to be, respectively, the CO<sub>2</sub> output (Fig. 7) and its uncertainty for each area.

## Results and discussion

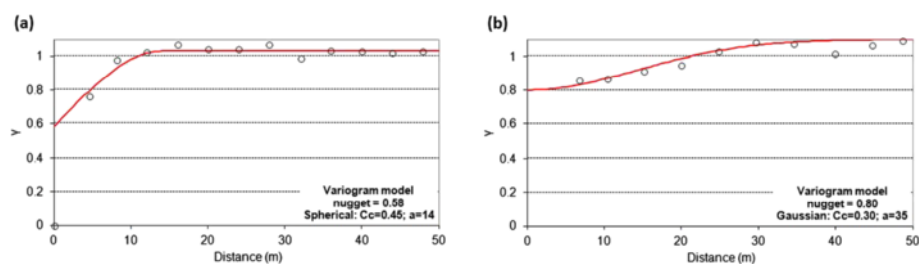
### TDL-based CO<sub>2</sub> distribution maps

Figure 4 is a distribution map of atmospheric CO<sub>2</sub> concentrations, calculated for a horizontal air cross-section taken at a 1.20-m height above the degassing soil and fumaroles of Furnas Lake (see Fig. 4). The calculated CO<sub>2</sub> concentration values range from ~390 to >850 ppm. In detail, the lowest CO<sub>2</sub> concentrations are detected in the eastern portion of the investigated area, far from the fumarolic field (close to laser positions B–E in Fig. 4). In contrast, two main clusters of higher CO<sub>2</sub> concentrations (>580) are detected on the northern, northwestern, and western portion of the map, in the proximity of the main degassing fumarolic vents (crosses). Peak CO<sub>2</sub> concentrations (>850 ppm), in particular, are observed close to the fumaroles WJ, MP, and ST (Fig. 4). It is worth noting that the CO<sub>2</sub> peaks appear in the map as shifted toward the east-southeast relative to the fumaroles' positions, in agreement with the prevalent plume transport direction during the observations (see wind rose diagram in Fig. 4).

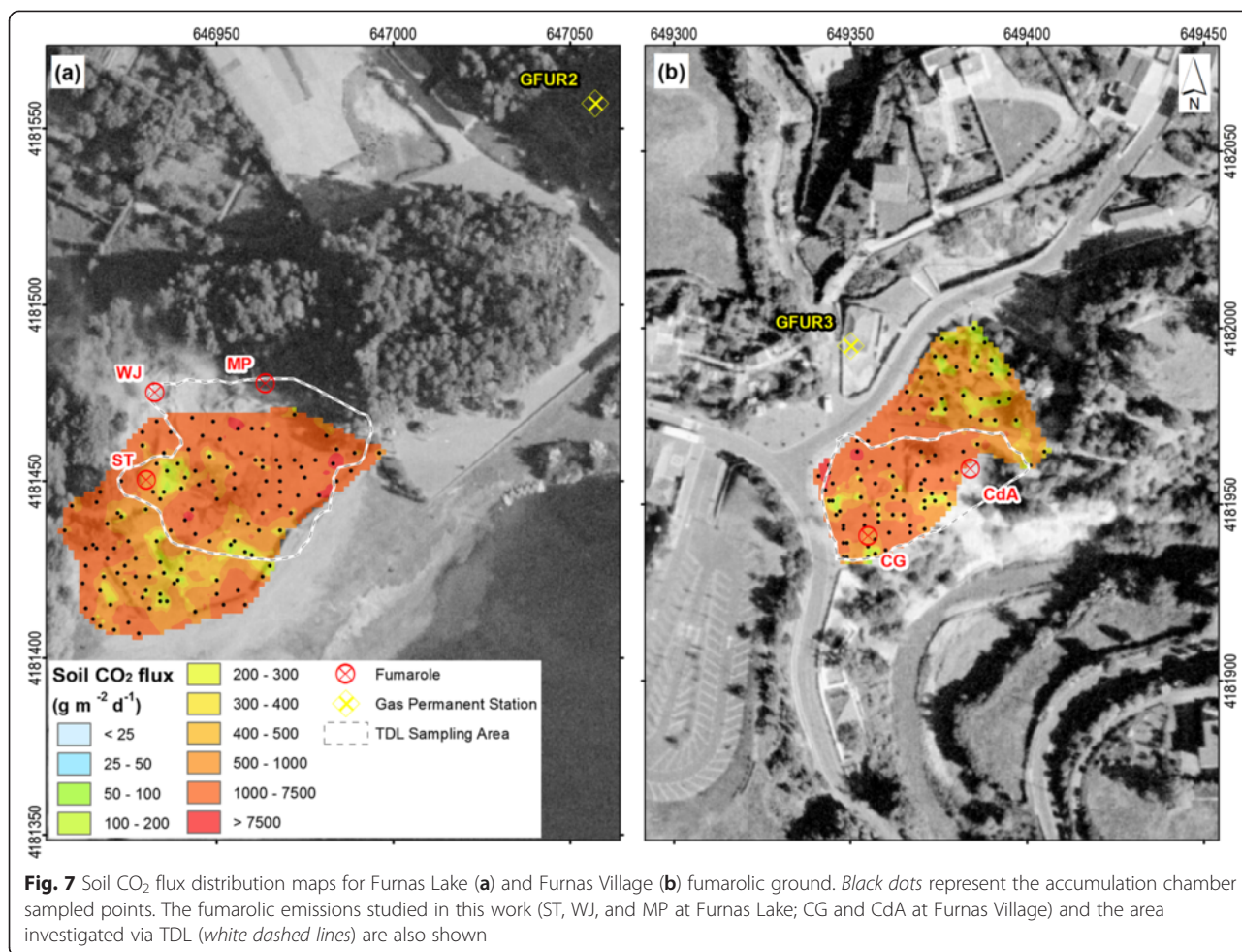
Figure 5 is a similar map derived from interpolation of the GasFinder dataset acquired during operation field at Furnas Village. Again, the CO<sub>2</sub> concentration anomalies are fairly consistent with the location of the main visible degassing areas. CO<sub>2</sub> concentrations (~700–800 ppm) peak in a wide degassing area downwind (west of) fumarole CdA (the dominant plume transport directions during the observations were toward NW and SW; Fig. 5). A secondary, more moderate CO<sub>2</sub> anomaly is observed north of CG fumarole. In contrast, the lowest CO<sub>2</sub> concentrations are detected along paths A4 and E6, respectively, upwind and/or more remote from the fumaroles (Fig. 5; Additional file 1: Table S2).

### Soil CO<sub>2</sub> flux degassing

Soil CO<sub>2</sub> fluxes (derived by using the accumulation chamber method) varied between 6 and 28,231 g m<sup>-2</sup> day<sup>-1</sup> in



**Fig. 6** Experimental and modeled variograms of the soil CO<sub>2</sub> flux normal scores. **a** Furnas Lake. **b** Furnas Village datasets. Legend: Cc (partial sill); a (range, m)



the Furnas Lake area, with a mean value of  $1510\text{ g m}^{-2}\text{ day}^{-1}$ . For the Furnas Village fumarolic ground, soil  $\text{CO}_2$  flux average value was  $1279\text{ g m}^{-2}\text{ day}^{-1}$ . Minimum and maximum values were, respectively, 5 and  $22,154\text{ g m}^{-2}\text{ day}^{-1}$  in this site (Table 2). Experimental variograms show spherical and Gaussian structures, respectively, for Furnas Lake and Furnas Village datasets (Fig. 6). The high nugget effect associated to Furnas Village (Fig. 6b) detailed survey is probably explained by the soil heterogeneities in this fumarolic ground, where the soil alteration causes large differences in the permeability at

small scales. This behavior was previously observed by Viveiros et al. (2010) for data acquired in the same sampling sites. Figure 7 shows the soil  $\text{CO}_2$  flux distribution maps for Furnas Lake and Furnas Village fumarolic ground.

**Multi-GAS in-plume measurements**

The Multi-GAS detected strong volcanic  $\text{CO}_2$  and  $\text{H}_2\text{S}$  signals in both areas.  $\text{SO}_2$  was undetected ( $<0.05\text{ ppmv}$ ) at both sites, and the relative humidity in the plumes ranged 30–50 % and 40–60 % at Furnas Lake and

**Table 2** Comparison between soil  $\text{CO}_2$  fluxes at Furnas Lake and Furnas Village during 19 and 22 August 2014 (this study) and soil  $\text{CO}_2$  fluxes calculated in previous study (March 2008 at Furnas Lake and June–July 2009 at Furnas Village). Areas (in  $\text{m}^2$ ), soil  $\text{CO}_2$  flux, mean, median, minimum, maximum values (expressed in  $\text{g m}^{-2}\text{ day}^{-1}$ ), and skewness are shown.  $S_p$  sampling period (year),  $N_p$  number of points

Site	$S_p$	$N_p$	Area	Mean	Median	Minimum	Maximum	Skewness
Lake	2014	124	3666	1510	266	6	28,231	5
Village	2014	95	2300	1279	248	5	22,154	4
Lake	2008	59	3666	947	187	6	20,579	6
Village	2009	87	2300	448	231	1	3452	3



Village, respectively. The relative humidity variations in the plumes did not systematically correlate with changes in either CO<sub>2</sub> or H<sub>2</sub>S, precluding the volcanic H<sub>2</sub>O signal to be clearly resolved.

An example of Multi-GAS acquisition is shown in Fig. 8, in which the CO<sub>2</sub> and H<sub>2</sub>S signals detected in 215 measurements close to “WJ” degassing vent were plotted. At Furnas Lake, the Multi-GAS derived compositions of the plumes of the main jet-water degassing vent (WJ; Figs. 2 and 4) are consistent with literature data based on direct fumarole sampling. We obtain a mean Multi-GAS-based CO<sub>2</sub>/H<sub>2</sub>S plume (molar) ratio of 353, which fits well with the mean ratio of 348 observed in an earlier chemical survey of the fumarole (Ferreira and Oskarsson 1999). At Furnas Village, the mean CO<sub>2</sub>/H<sub>2</sub>S plume ratio of 150 measured by the Multi-GAS at the CG fumarole (Figs. 3 and 5) is not far from the mean fumarolic ratio of 120 of Ferreira and Oskarsson (1999). The CO<sub>2</sub>/S<sub>t</sub> ratios of Furnas fumaroles (this study) are similar to those seen in other hydrothermal fluids of similar outlet temperatures. For example, the CO<sub>2</sub>/S<sub>t</sub> ratio detected at Furnas Lake is close to that quoted by Aiuppa et al. (2013) for Pisciarelli fumaroles of Campi Flegrei (CO<sub>2</sub>/S<sub>t</sub> ratio of 300, range 210–410) and by

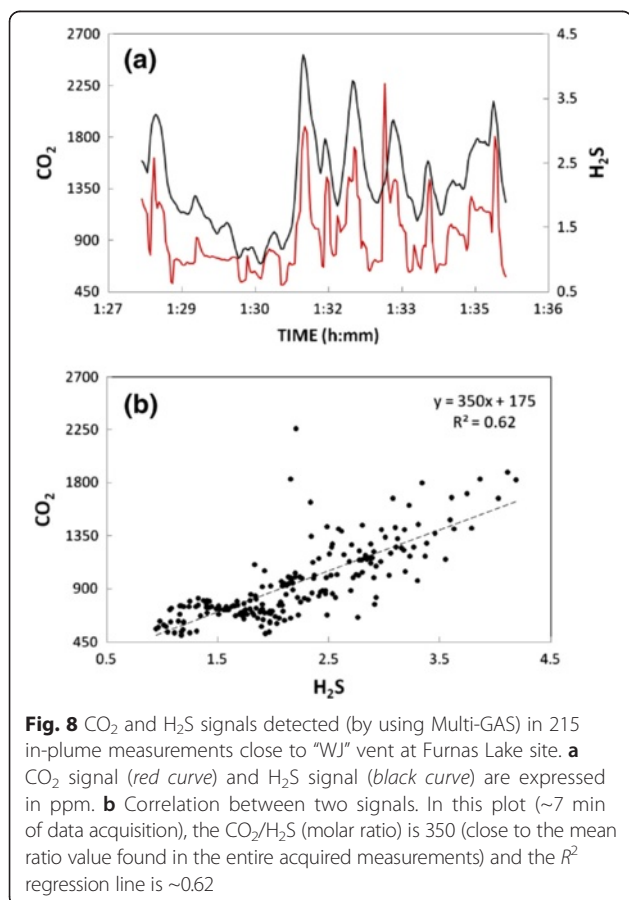
Chiodini et al. (2001b) for Vesuvius fumaroles (CO<sub>2</sub>/S<sub>t</sub> ratio of 342). The CO<sub>2</sub>/S<sub>t</sub> ratio detected at Furnas Village is comparable with compositions of fumaroles at Solfatara crater, Campi Flegrei (CO<sub>2</sub>/S<sub>t</sub> ratio of 130, range 120–152; Aiuppa et al. 2013) and Ischia (CO<sub>2</sub>/S<sub>t</sub> ratio of 119; Chiodini et al. 2004).

#### CO<sub>2</sub> and H<sub>2</sub>S output

We estimated the fumarolic CO<sub>2</sub> output from Furnas volcano by 2D integration of the TDL-based CO<sub>2</sub> concentration values. This integration was restricted to the shaded box areas of Figs. 4 and 5, downwind the main fumaroles. These integrated CO<sub>2</sub> amounts within the defined areas were finally multiplied by the vertical gas transport speeds to calculate the fluxes (Table 1). The so-estimated fumarolic CO<sub>2</sub> fluxes were ~29 t day<sup>-1</sup> at Furnas Lake (17.6 ± 5.3 t day<sup>-1</sup> emitted from WJ and 11.4 ± 4.7 t day<sup>-1</sup> from ST) and ~21 t day<sup>-1</sup> at Furnas Village (17.8 ± 4.1 t day<sup>-1</sup> emitted from CdA and 3 ± 0.8 t day<sup>-1</sup> from CG) (Table 1). The cumulative fumarolic output from the two areas is therefore ~50 t day<sup>-1</sup>.

We additionally infer, using this fumarolic CO<sub>2</sub> flux (Table 1) and the Multi-GAS-derived CO<sub>2</sub>/H<sub>2</sub>S ratios, fumarolic H<sub>2</sub>S fluxes of 0.08 t day<sup>-1</sup> (Furnas Lake) to 0.17 t day<sup>-1</sup> (Furnas Village), respectively. Our results indicate that Furnas fumaroles are weak, but yet not negligible, sources of S. The present-day H<sub>2</sub>S flux of 0.25 t day<sup>-1</sup> (0.08 + 0.17 t day<sup>-1</sup>) at Furnas is nearly 1 order of magnitude lower than typical emissions from Campi Flegrei volcano (1.5–2.2 t day<sup>-1</sup>; Aiuppa et al. 2013), which similarly display a H<sub>2</sub>S-dominated flux and more than 1 order of magnitude lower than emissions from Vulcano Island (6–9 t day<sup>-1</sup>; Tamburello et al. 2011), whose fumarolic field hosts hot (>400 °C) magmatic fumaroles. For comparison, the H<sub>2</sub>S flux sustained by Etna, the largest source of volcanic gases worldwide, is 200 to 400 times larger (50 to 113 t day<sup>-1</sup>; Aiuppa et al. 2005) (note Etna additionally emits thousands of tons of SO<sub>2</sub> every day; Caltabiano et al. 2004).

On the basis of the integration of the average of 100 sequential Gaussian simulations (Deutsch and Journel 1998; Cardellini et al. 2003), over the sampled area, mean soil diffuse CO<sub>2</sub> emissions of ~6 ± 0.2 t day<sup>-1</sup> and 3.2 ± 0.2 t day<sup>-1</sup> were estimated for Furnas Lake and Furnas Village degassing areas, respectively. We then obtain a cumulative soil CO<sub>2</sub> release of ~9.2 t day<sup>-1</sup>. Detailed soil CO<sub>2</sub> surveys in the Furnas Lake and Furnas Village fumarolic fields (same as Fig. 7 in this study) have previously been performed by Viveiros et al. (2012). For the sake of comparison, we extracted from these earlier surveys (made between 2008 and 2009) the results obtained in the same areas as those investigated in the present study (2014) (Fig. 7). This comparison shows that the Furnas Lake area emitted in 2008 about 3 t day<sup>-1</sup> of



CO<sub>2</sub>, or about half of what was measured in 2014. Similarly, the Furnas Village area was found in 2009 to emit nearly half (1.7 t day<sup>-1</sup>) of the amount of CO<sub>2</sub> released in 2014 (3.2 ± 0.2 t day<sup>-1</sup>). Increased CO<sub>2</sub> emissivity in 2014 is consistent with the different average CO<sub>2</sub> flux values observed in the two different surveys (Table 2) but is not supported by records of the permanent soil CO<sub>2</sub> flux station GFUR2 (where lower soil CO<sub>2</sub> flux values—about 250 g m<sup>-2</sup> day<sup>-1</sup>—were recorded during 2014 compared to 2008 and 2009 surveys—values higher than 350 g m<sup>-2</sup> day<sup>-1</sup>). We caution that differences in sampling grid/density (that was significantly higher in the present surveys, as can be also observed in the number of points of Table 2) can partially explain the CO<sub>2</sub> flux diversity between the two campaigns (Viveiros et al. 2010).

### Implications

Our results here suggest that, in the actively degassing fumarolic areas of Furnas, fumarolic vents, with their cumulative fumarolic output of ~50 t day<sup>-1</sup>, dominate the total CO<sub>2</sub> degassing budget (~59 t day<sup>-1</sup>) and overwhelm the relatively marginal contribution (~15 %) of soil diffuse degassing. While the fumarolic output can be locally important, however, its contribution to the total CO<sub>2</sub> degassing output becomes marginal at the scale of the entire volcano. Earlier soil CO<sub>2</sub> flux measurements, in fact, have demonstrated a total diffusive hydrothermal CO<sub>2</sub> release at Furnas, from an area (6.2 km<sup>2</sup>) far larger than that studied here, as high as 968 t day<sup>-1</sup> (Viveiros et al. 2010). From this comparison, we conclude that—at Furnas—the most actively degassing areas, although featuring the most visible (e.g., fumaroles, hot pools) and spectacular manifestations of thermal activity, in no way correspond to the areas of largest CO<sub>2</sub> output: in total, the actively degassing areas contribute only about 5 % to the total CO<sub>2</sub> output.

To what extent this conclusion can be generalized to other hydrothermal-volcanic systems remains unknown, given the paucity of information we have in hand. A small contribution of fumarolic gas vents to the total CO<sub>2</sub> budget was also suggested by Fridriksson et al. (2006) for the Reykjanes geothermal area, SW Iceland, where the contribution of steam vents/mud pools is only about 3 % of the soil CO<sub>2</sub> degassing output. In other systems, in contrast, the contribution of fumaroles is manifestly more significant. At Mud volcano (Yellowstone, USA), for example, Werner et al. (2000) estimated that fumarolic vent emissions contribute to more than 32 % of the total degassing; based on statistical approach, the authors even suggested that the hydrothermal-focused emissions can be responsible for up to 63 % of the total degassing in that thermal area. At Campi Flegrei, the CO<sub>2</sub> fumarolic output of ~500 t day<sup>-1</sup> (range, 460–507 t day<sup>-1</sup>, Aiuppa et al. 2013; Pedone et al. 2014a) makes a

substantial contribution to the total CO<sub>2</sub> output, which is still dominated by soil diffuse degassing (~1100 ± 120 t day<sup>-1</sup>; Chiodini et al. 2010). From these examples, we argue that much remains to be done to fully understand the fumarolic vs. diffusive gas contribution at volcano-hydrothermal systems. The relative significance of the fumarolic output can differ significantly from one volcano to another. A general conclusion is that the relative significance of the above two forms of gas dissipation will likely be dependent upon the scale at which the comparison is made. At the local scale of an active hydrothermal manifestation, the fumarolic output will likely overwhelm the soil diffuse flux, while the latter will most likely dominate at a larger scale (at the scale of an entire volcanic complex). For example, based on results from 20 analyzed hydrothermal areas, Harvey et al. (2015) concluded that the contribution of focused venting to the total CO<sub>2</sub> emission is typically less than 10 %. The TDL, with the ability to characterize the fumarolic CO<sub>2</sub> output, promises to contribute a substantial advancement in this field in the following years.

### Conclusions

We estimated the total CO<sub>2</sub> emissions from the main thermal manifestations of Furnas volcano by jointly using two different techniques: the GasFinder 2.0 tunable diode laser and the accumulation chamber method. We find that, in the most vigorously degassing areas, the soil CO<sub>2</sub> flux contribution (approximately 9.2 t day<sup>-1</sup>) represents a minor (~18 %) contribution to the total CO<sub>2</sub> output, which is dominated by the fumaroles (about ~50 t day<sup>-1</sup>). The CO<sub>2</sub> output contributed by the fumaroles is larger than that contributed by Furnas springs (~12 t day<sup>-1</sup>, Cruz et al. 1999), but far lower than the total hydrothermal diffuse degassing flux (~968 t day<sup>-1</sup>) at the scale of the entire volcano. This observation supports the conclusions that although fumaroles are the most visible surface manifestations of thermal activity, they are not necessarily the biggest contributors to the total CO<sub>2</sub> output from quiescent, Solfataro-stage volcanoes, where CO<sub>2</sub> is mainly released in silent, invisible form through soil emissions.

Summing up the CO<sub>2</sub> flux contributions for the fumarolic emissions, the degassing soils and the springs (12 t day<sup>-1</sup>, Cruz et al. 1999), the total volcanic/hydrothermal CO<sub>2</sub> output for Furnas volcano is estimated to be ~1030 t day<sup>-1</sup>. These results show once more the importance of taking into consideration both soil degassing and gas vent emissions to estimate the CO<sub>2</sub> emission in hydrothermal areas, since their relative contribution seems to be quite different depending on the study sites. Using a portable Multi-GAS, we also obtained the CO<sub>2</sub>/H<sub>2</sub>S ratio signature for the investigated fumaroles, and concluded that the Furnas fumaroles are, in their present

state of activity, weak, but yet not negligible, sources of H<sub>2</sub>S (~0.25 t day<sup>-1</sup>).

## Additional file

### Additional file 1: The supplementary file is a more detailed documentation about the TDL acquisitions and data elaboration.

A1. CO<sub>2</sub> TDL-datasets. A2. Parameters used to perform sGs with CO<sub>2</sub> concentrations TDL data and Zonal Statistic on E-Type maps. A3. Parameters used to perform sGs with soil CO<sub>2</sub> flux data (accumulation chamber). In A1 section, each path laser-retroreflector acquisition during the campaigns carried out at Furnas Lake (Additional file 1: Table S1) and Furnas Village (Additional file 1: Table S2) is shown. In A2 section, more details about statistical approach and elaboration of data to create the distribution CO<sub>2</sub> concentration maps are shown (Additional file 1: Tables S3, S4, and S5). Finally, parameters used to perform sGs with soil CO<sub>2</sub> flux data are shown in A3 section (Tables S6 and S7).

## Competing interests

The authors declare that they have no competing interests.

## Authors' contributions

MP and FV carried out the field campaigns and drafted the manuscript. AA allowed the work realization and actively contributed to drafting the manuscript. GG, FG, and TF provided important suggestions during data processing. FV and ALG provided useful suggestions/help in the maps elaboration. VF provided technical assistance in laboratory. All authors have read and approved the final manuscript.

## Authors' information

MP is a Dr. in geochemistry at DiSTeM (Dipartimento di Scienze della Terra e del Mare) of Palermo University and has a study grant at INGV (Istituto Nazionale di Geofisica e Vulcanologia) of Palermo. FV has a post-doc grant from Fundação para a Ciência e Tecnologia (FCT) and is a researcher at Centro de Vulcanologia e Avaliação de Riscos Geológicos (CVARG) of the University of the Azores. AA is Associate Professor of Volcanology and Geochemistry at DiSTeM (Dipartimento di Scienze della Terra e del Mare) of Palermo University and Associate Researcher at Istituto Nazionale di Geofisica e Vulcanologia (Sezione di Palermo). GG is first engineer at Istituto Nazionale di Geofisica e Vulcanologia (Sezione di Palermo). FG is a researcher at Istituto Nazionale di Geofisica e Vulcanologia (Sezione di Palermo). ALG is a Dr. in geochemistry at DiSTeM (Dipartimento di Scienze della Terra e del Mare) of Palermo University and has a study grant at INGV of Palermo. VF is a technician at Istituto Nazionale di Geofisica e Vulcanologia (Sezione di Palermo). TF is Associate Professor in the Department of Geosciences (University of the Azores) and Researcher at Centro de Vulcanologia e Avaliação de Riscos Geológicos (CVARG).

## Acknowledgements

The research leading to these results has received funding from the European Research Council under the European Union's Seventh Framework Programme (FP7/2007/2013)/ERC grant agreement n° 305377, and from the FP7 grant "Futurevolc." F. Viveiros is supported by a post-doctoral grant from Fundação para a Ciência e a Tecnologia (FCT). The authors would like to acknowledge the technical assistance from the staff of Centro de Vulcanologia e Avaliação de Riscos Geológicos, University of the Azores, and the staff of INGV (Istituto Nazionale di Geofisica e Vulcanologia) of Palermo, Italy. The authors would also like to thank the two reviewers, Dr. C. Werner and Dr. H. Shinohara, for the valuable contribution to improve the early version of the manuscript.

## Author details

<sup>1</sup>DiSTeM, Università di Palermo, via Archirafi, 36, Palermo 90123, Italy. <sup>2</sup>Istituto Nazionale di Geofisica e Vulcanologia, Sezione di Palermo, via Ugo La Malfa, 153, Palermo 90146, Italy. <sup>3</sup>Centro de Vulcanologia e Avaliação de Riscos Geológicos, University of the Azores, Rua Mãe de Deus, Ponta Delgada 9501-801, Portugal.

Received: 24 June 2015 Accepted: 15 October 2015

Published online: 26 October 2015

## References

- Aiuppa A, Inguaggiato S, McGonigle AJS, O'Dwyer M, Oppenheimer C, Padgett MJ, Rouwet D, Valenza M (2005) H<sub>2</sub>S fluxes from Mt. Etna, Stromboli and Vulcano (Italy) and implications for the global volcanic sulfur budget. *Geochim Cosmochim Acta* 69(7):1861–1871. doi:10.1016/j.gca.2004.09.018
- Aiuppa A, Burton M, Allard P, Caltabiano T, Giudice G, Gurrieri S, Liuzzo M, Salerno G (2011) First observational evidence for the CO<sub>2</sub>-driven origin of Stromboli's major explosions. *Solid Earth* 2:135–142
- Aiuppa A, Giudice G, Liuzzo M, Tamburello G, Allard P, Calabrese S, Chaplygin I, McGonigle AJS, Taran Y (2012) First volatile inventory for Gorely volcano, Kamchatka. *Geophys Res Lett* 39:L06307. doi:10.1029/2012GL051177
- Aiuppa A, Tamburello G, Di Napoli R, Cardellini C, Chiodini G, Giudice G, Grassa F, Pedone M (2013) First observations of the fumarolic gas output from a restless caldera: implications for the current period of unrest (2005–2013) at Campi Flegrei. *Geochemistry Geophysics Geosystems*. doi:10.1002/ggge.20261
- Baubron JC, Baxter P, Coutinho R, Allard P, Ferreira T, Gaspar JL (1994) Methodology for the drawing up of the Furnas gas hazard map. In: Barberi F, Casale R, Fratta M (eds) *The European laboratory volcanoes. Workshop Proceedings*, Catania, pp 262–267
- Baxter P, Baubron JC, Coutinho R (1999) Health hazards and disaster potential of ground gas emissions at Furnas Volcano, São Miguel, Azores. *J Volcanol Geotherm Res* 92:95–106
- Caliro S, Viveiros F, Chiodini G, Ferreira T (2015) Gas geochemistry of hydrothermal fluids of the S. Miguel and Terceira Islands, Azores. *Geochim Cosmochim Acta*. doi:10.1016/j.gca.2015.07.009
- Caltabiano T, Burton M, Giammanco S, Allard P, Bruno N, Muré F, Romano R (2004) Volcanic gas emissions from Mt. Etna, 1987–2000. In: Bonaccorso A, Calvari S, Coltelli M, Del Negro C, Falsaperla S (eds) *Mt. Etna: Volcano Laboratory*, vol. 143, *Geophys Monogr Ser. AGU*, Washington, DC, pp 111–128
- Cardellini C, Chiodini G, Frondini F (2003) Application of stochastic simulation to CO<sub>2</sub> flux from soil: mapping and quantification of gas release. *J Geophys Res* 108(B9):2425–2437. doi:10.1029/2002JB002165
- Chiodini C, Cioni R, Guidi M, Raco B, Marini L (1998) Soil CO<sub>2</sub> flux measurements in volcanic and geothermal areas. *Appl Geochem* 13:543–552
- Chiodini G, Frondini F, Kerrick DM, Rogie J, Parello F, Peruzzi L, Zanzari AR (1999) Quantification of deep CO<sub>2</sub> fluxes from Central Italy. Examples of carbon balance for regional aquifers and of soil diffuse degassing. *Chem Geol* 159:205–222
- Chiodini G, Frondini F, Cardellini C, Granieri D, Marini L, Ventura G (2001a) CO<sub>2</sub> degassing and energy release at Solfatara Volcano, Campi Flegrei, Italy. *J Geophys Res* 106:16213–16221
- Chiodini G, Marini L, Russo M (2001b) Geochemical evidence for the existence of high-temperature hydrothermal brines at Vesuvio Volcano, Italy. *Geochim Cosmochim Acta* 65(13):2129–2147
- Chiodini G, Avino R, Brombach T, Caliro S, Cardellini C, De Vita S, Frondini F, Granieri D, Marotta E, Ventura G (2004) Fumarolic and diffuse soil degassing west of Mount Epomeo, Ischia, Italy. *J Volcanol Geotherm Res* 133(1–4):291–309. doi:10.1016/j.jvolgeores.2003.02.00
- Chiodini G, Caliro S, Cardellini C, Granieri D, Avino R, Baldini A, Donnini M, Minopoli C (2010) Long-term variations of the Campi Flegrei, Italy, volcanic system as revealed by the monitoring of hydrothermal activity. *J Geophys Res* 115:B03205. doi:10.1029/2008JB006258
- Cole P, Queiroz G, Wallenstein N, Gaspar JL, Duncan AM, Guest J (1995) An historic subplinian/phreatomagmatic eruption: the 1630 AD eruption of Furnas Volcano, São Miguel, Azores. *J Volcanol Geotherm Res* 69:117–135
- Cruz JV, Coutinho R, Carvalho MR, Óskarsson N, Gislason SR (1999) Chemistry of waters from Furnas Volcano, São Miguel, Azores: fluxes of volcanic carbon dioxide and leached material. *J Volcanol Geotherm Res* 92:151–167
- De Natale P, Gianfrani L, De Natale G, Cioni R (1998) Gas concentration measurements with DFB lasers to monitor volcanic activity. *SPIE Proceedings Series. Applications of Photonic Technology-3* 3491:783–787
- Deutsch CV, Journel AG (1998) *GSLIB: Geostatistical Software Library and user's guide*. 369pp, Applied Geostatistics Series. Oxford University Press, New York
- Ferreira T, Óskarsson N (1999) Chemistry and isotopic composition of fumarole discharges of Furnas caldera. *J Volcanol Geotherm Res* 92:169–179
- Ferreira T, Gaspar JL, Viveiros F, Marcos M, Faria C, Sousa F (2005) Monitoring of fumarole discharge and CO<sub>2</sub> soil degassing in the Azores: contribution to volcanic surveillance and public health risk assessment. *Ann Geophys* 48(4–5):787–796
- Fridriksson T, Kristjánsson BR, Ármannsson H, Margrétardóttir E, Ólafsdóttir S, Chiodini G (2006) CO<sub>2</sub> emissions and heat flow through soil, fumaroles, and

- steam heated mud pools at the Reykjanes geothermal area, SW Iceland. *Appl Geochem* 21:1551–1569
- Gianfrani L, Gagliardi G, Pesce G, Sasso A (1997a) High sensitivity detection of NO<sub>2</sub> by using a 740 nm semiconductor diode laser. *Appl Phys B* 64:487
- Gianfrani L, Gabrysch M, Corsi C, De Natale P (1997b) Detection of H<sub>2</sub>O and CO<sub>2</sub> with distributed feedback diode lasers: measurement of broadening coefficients and assessment of the accuracy levels for volcanic monitoring. *Appl Opt* 36:9481–9486
- Gianfrani L, De Natale P, De Natale G (2000) Remote sensing of volcanic gases with a DFB-laser-based fiber spectrometer. *Appl Phys B-Rapid Commun* 70:467–470
- Harvey MC, Rowland JV, Chiodini G, Rissmann CF, Bloomberg S, Hernández PA, Mazot A, Viveiros F, Werner C (2015) Heat flux from magmatic hydrothermal systems related to availability of fluid recharge. *J Volcanol Geotherm Res* 302:225–236
- Hernández PA, Salazar JM, Shimoike Y, Mori T, Notsu K, Perez N (2001) Diffuse emission of CO<sub>2</sub> from Miyakejima volcano, Japan. *Chem Geol* 177:175–185
- Moore R (1990) Volcanic geology and eruption frequency, São Miguel, Azores. *Bull Volcanol* 52:602–614
- Pedone M, Aiuppa A, Giudice G, Grassa F, Cardellini C, Chiodini G, Valenza M (2014a) Volcanic CO<sub>2</sub> flux measurement at Campi Flegrei by tunable diode laser absorption spectroscopy. *Bull Volcanol* 76. doi:10.1007/s00445-014-0812-z
- Pedone M, Aiuppa A, Giudice G, Grassa F, Francofonte V, Bergsson B, Ilyinskaya E (2014b) Tunable diode laser measurements of hydrothermal/volcanic CO<sub>2</sub>, and implications for the global CO<sub>2</sub> budget. *Solid Earth* 5:1209–1221. doi:10.5194/sed-5-1209-2014
- Richter D, Erdelyi M, Curl RF, Tittel FK, Oppenheimer C, Duffell HJ, Burton M (2002) Field measurements of volcanic gases using tunable diode laser based mid-infrared and Fourier transform infrared spectrometers. *Opt Lasers Eng* 37:171–186
- Rogie JD, Kerrick DM, Sorey ML, Chiodini G, Galloway DL (2001) Dynamics of carbon dioxide emission at Mammoth Mountain, California. *Earth Planet Sci Lett* 188:535–541
- Searle R (1980) Tectonic pattern of the Azores spreading centre and triple junction. *Earth Planet Sci Lett* 51:415–434
- Tamburello G, Kantzas EP, McGonigle AJS, Aiuppa A, Giudice G (2011) UV camera measurements of fumarole field degassing (La Fossa crater, Vulcano Island). *J Volcanol Geotherm Res* 199:47–52
- Tulip J (1997) Gas detector, United States Patent 5,637,872 250/338.5
- Viveiros F, Ferreira T, Cabral Vieira J, Silva C, Gaspar JL (2008) Environmental influences on soil CO<sub>2</sub> degassing at Furnas and Fogo volcanoes (São Miguel Island, Azores archipelago). *J Volcanol Geotherm Res* 177:883–893
- Viveiros F, Cardellini C, Ferreira T, Caliro S, Chiodini G, Silva C (2010) Soil CO<sub>2</sub> emissions at Furnas volcano, São Miguel Island, Azores archipelago: volcano monitoring perspectives, geomorphologic studies, and land use planning application. *J Geophys Res* 115:B12208. doi:10.1029/2010JB007555
- Viveiros F, Cardellini C, Ferreira T, Silva C (2012) Contribution of CO<sub>2</sub> emitted to the atmosphere by diffuse degassing from volcanoes: the Furnas Volcano case study. *Int J Global Warm* 4(3–4):287–304
- Werner C, Brantley SL, Boomer K (2000) CO<sub>2</sub> emissions related to the Yellowstone volcanic system. Statistical sampling, total degassing, and transport mechanisms. *J Geophys Res* 105:10831–10846
- Werner C, Hurwitz S, Evans WC, Lowenstern JB, Bergfeld D, Heasler H, Jaworowski C, Hunt A (2008) Volatile emissions and gas geochemistry of Hot Spring Basin, Yellowstone National Park, USA. *J Volcanol Geotherm Res* 178:751–762

Submit your manuscript to a SpringerOpen® journal and benefit from:

- Convenient online submission
- Rigorous peer review
- Immediate publication on acceptance
- Open access: articles freely available online
- High visibility within the field
- Retaining the copyright to your article

---

Submit your next manuscript at ► [springeropen.com](http://springeropen.com)

---